

**INTRODUCTION**

The Chatham 7.5-minute topographic quadrangle lies within the Piedmont Physiographic Province in north-central New Jersey. Somerset and Union Counties cover the southeastern half of the quadrangle and Morris County occupies the northwestern half. Development has been confined to the valleys and higher, more level topography. Recently, developers have begun the push towards the ridge slopes and create Interstate Route 78 provides easy access to industry and business centers to the east. The Great Swamp National Wildlife Refuge covering a large area in the north center of the quadrangle remains a large wildlife habitat amongst the developed regions of northern New Jersey.

The surficial and bedrock geology controls the varied topography in the quadrangle. Ridges underlain by igneous units delineate a strong northeast trend with elevations ranging from 400 to 550 feet above mean sea level. Sedimentary rocks mark the lower elevations from 90 to 220 feet between the igneous ridges. Glacial sediments that mute the bedrock geology form both the flat terrain of the Great Swamp and the slightly hummocky northwest trending topography northeast of the Great Swamp. Recent fluvial deposits from modern rivers are the latest features that shaped the current landscape of the quadrangle. Of these rivers the Passaic River is the largest and bisects the quadrangle along a northeast trend.

**STRATIGRAPHY**

Surficial cover dominates the geology of the Great Swamp National Wildlife refuge and the region to the northeast. These sediments consist of Pleistocene glacial units and Quaternary alluvial deposits. Glacial lake-bottom sediments control the flat lying Great Swamp topography (Salisbury and Kummel, 1959). The varied topography in the towns of Chatham and Madison marks the presence of the late Pleistocene terminal moraine (Stanford, unpublished data). Recent alluvial deposits mirror the trends of streams and rivers. A thin veneer of soil blankets the bedrock in the higher areas.

The bedrock units in this region developed in the Mesozoic during the initial breakup of the supercontinent of Pangea. Southeast-dipping normal faults controlled the formation of the Newark rift basin on the east. Continued episodic motion on these faults controlled the sediment input into the depression, both from the Highlands to the west and from an eastern source. Intra-basinal faulting was also active during deposition (Schische, 1992, 1993). Bedrock units range from Early Jurassic to Late Triassic in age (Olsen, 1980a). They consist of a series of three basalt units separated by interbedded fluvial and lacustrine sediments. Each of the basalt units, from youngest to oldest the Hook Mountain, Preakness and Orange Mountain, consists of multiple flows. Van Houten (1969) described the internal structure common to many of the flows containing a lower and upper colonnade separated by an entablature zone. A basal massive basalt grading upward through a platy zone and into regular columns forms the lower colonnade. The entablature is the core of the flow and contains curvilinear columnar structures (Van Houten, 1969; Faust, 1978). The upper colonnade is marked by a lower pseudo-columnar zone capped by massive basalt and locally scoria. The development of each of the three parts varies within the flows (Olsen, 1980a; Olsen and others, 1989). Pillow lavas are present locally (Van Houten, 1969). A thin sedimentary unit is mapped as part of the Preakness Basalt. Exposure of this unit in the quadrangle is sparse. Field notes from Henry Kummel (Kummel, 1900) located one area of outcrop where float chips of red-brown siltstone occur. A second locality within a new house foundation aided in delimiting the trend of the sediments. Evidence of the sedimentary layers also occurs in domestic well cuttings but its exact location in the eastern half of the quadrangle is problematic from the data patterns during this study.

The sedimentary units between the basalt formations portray a fluvial-to-lacustrine-deposition cyclicity from oldest to youngest; they consist of the Passaic, Felville, Towaco and Bonton formations. Units incorporate alternating, dominantly red-brown fluvial sediments, and gray to olive lacustrine deposits (Van Houten, 1969; Olsen, 1980a, 1980b; Olsen and others 1989). Fluvially deposited, sandstone through heavy siltstone to mudstone dominates over the siltstone to mudstone origin. Fedich and Smoot (1988) suggested a lower fluvial gradient and deeper lake levels from the Passaic through the Felville Formations deposition. Drier conditions prevailed during the Towaco and Bonton Formation deposition. Olsen and others (1996) show Milankovitch orbital cyclicity controlling the fluvial- and lacustrine-deposition. Gray bed cycles became important markers used in defining the various members of the sedimentary formations.

**STRUCTURE**

The Watchung syncline and Mount Vernon anticline, first outlined by Rogers (1840) are the major fold structures in the mapped area. Darton (1890) first used the terms "Watchung flexure" and subsequently, Watchung syncline. Both terms of the Watchung syncline dip gently, ranging from 0° to 15°. This produces a gentle broad open synclinal structure trending northeasterly. Pleistocene sediments blanket the fold axis but the trend of the axis parallels the trace of the Watchung Fault in the region. The north-south-trending Mount Vernon anticline projects from the Morristown quadrangle into the northwest corner of the study area (Ratcliffe and others, 1990). Fold limbs of the anticline also dip at gentle angles, producing a gentle open fold.

A series of north-trending normal faults, marking the termination of the New Brunswick Fault, cuts the Orange Mountain and Preakness Basalts and Felville Formations. The New Brunswick Fault traverses the Plainfield and New Brunswick quadrangles before being displaced and splaying out into a myriad of smaller faults. Excellent examples of the normal faults occur within the Veldon Quarry. Limited subsidiary normal faults can be found along the Passaic Formation/Orange Mountain Basalt front.

Figure 1 shows strike of bedding and fracture trends mapped and analyzed in 1999. Bedding partings were not classified as fractures but used to delineate strike of sedimentary units. Fractures within the sedimentary and igneous units cluster around due North and are not parallel to bedding trends.

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**DESCRIPTION OF MAP UNITS**

- Jb** Bonton Formation (Lower Jurassic) (Olsen, 1980a) - Reddish-brown to brownish-purple, fine-grained, commonly micaceous sandstone, siltstone, and mudstone, in fining-upward sequences mostly 5 to 13 ft thick. Red, gray, and brownish-purple siltstone and black, blocky, partly dolomitic siltstone and shale are common in the lower part of unit. Irregular mudcracks, symmetrical ripple marks, hummocky and trough cross-laminated, burrows, and evaporite minerals are abundant in red siltstone and mudstone. Gray, fine-grained sandstone may have carbonized plant remains and reptile footprints in middle and upper parts of unit. Maximum thickness regionally is about 1,640 ft.
- Jh** Hook Mountain Basalt (Lower Jurassic) (Olsen, 1980a) - Dark-greenish-gray to black, generally fine-grained and very locally medium- to coarse-grained, amygdaloidal basalt composed of plagioclase, clinopyroxene, and iron-titanium oxides. Contains small to large vesicles lined with prehnite and spheral to tubular gas-escape vesicles, some filled by zeolite minerals or calcite, near the base of flows. Dark-gray, coarse- to very-coarse-grained gabbroid (Jgp) composed of clinopyroxene grains as much as 0.5 ft. long and plagioclase grains as much as 1.0 ft. long occurs at several stratigraphic intervals in the unit. Upper contacts are sharp and lower contacts are gradational with more typical fine-grained basalt. Unit consists of at least three major flows, the tops of which are marked by prominent vesiculated zones as much as 8 ft. thick. Radiating slender columns 2 to 24 in. wide, due to shrinkage during cooling, are abundant near the base of the lowest flow. A thin, 6 to 25 ft., thick bed of reddish-brown siltstone (Jps) separates the lower flows. Maximum thickness of unit is about 360 ft.
- Jt** Towaco Formation (Lower Jurassic) (Olsen, 1980a) - Reddish-brown to brownish-purple, buff, olive-tan, or light-olive-gray, fine- to medium-grained, micaceous sandstone, siltstone, and siltly mudstone in fining-upward sequences 3 to 10 ft. thick. Unit consists of at least eight sequences of gray, greenish-gray, or brownish-gray, fine-grained sandstone, siltstone, and calcareous siltstone, and black microlaminated calcareous siltstone and mudstone with diagnostic pellet grains, dinosaur tracks and fish fossils, irregular mudcracks, symmetrical ripple marks present. Sandstone is commonly hummocky and trough cross-laminated and siltstone is commonly planar-laminated or bioturbated and indistinctly laminated to massive. Several inches of unit have been thermally metamorphosed along contact with Hook Mountain Basalt. Maximum thickness is about 1,250 ft.
- Jp** Preakness Basalt (Lower Jurassic) (Olsen, 1980a) - Dark-greenish-gray to black, fine-grained, dense, hard basalt composed mainly of intergrown plagioclase and clinopyroxene. Contains small tubular gas-escape vesicles, some filled by zeolite minerals or calcite, near the base of flows. Dark-gray, coarse- to very-coarse-grained gabbroid (Jgp) composed of clinopyroxene grains as much as 0.5 ft. long and plagioclase grains as much as 1.0 ft. long occurs at several stratigraphic intervals in the unit. Upper contacts are sharp and lower contacts are gradational with more typical fine-grained basalt. Unit consists of at least three major flows, the tops of which are marked by prominent vesiculated zones as much as 8 ft. thick. Radiating slender columns 2 to 24 in. wide, due to shrinkage during cooling, are abundant near the base of the lowest flow. A thin, 6 to 25 ft., thick bed of reddish-brown siltstone (Jps) separates the lower flows. Maximum thickness of unit is about 1,040 ft.
- Jps** Felville Formation (Lower Jurassic) (Olsen, 1980a) - Reddish-brown, or light-grayish-red, fine- to coarse-grained sandstone, siltstone, shaly siltstone, and siltly mudstone, and light- to dark-gray or black, locally calcareous siltstone, siltly mudstone, and carbonaceous limestone. Upper part of unit is predominantly thin- to medium-bedded, reddish-brown siltstone. Reddish-brown sandstone and siltstone are moderately well sorted, commonly cross-laminated, and interbedded with reddish-brown, planar-laminated siltly mudstone to mudstone. Two thin, laterally continuous sequences, each as much as 10 ft. thick, consisting of dark-gray to black, carbonaceous limestone, light-gray limestone, and medium-gray calcareous siltstone, shale or clay, desiccated shale to silty shale occur near the base, and together with the red beds below, comprise the Washington Valley Member (Jw) of Olsen (1980b). Gray beds may contain fish, reptiles, arthropods, and diagnostic plant fossils. Several inches of unit have been thermally metamorphosed along contact with Preakness Basalt (Jp). Thickness ranges from 450 to 465 ft.

- Jo** Orange Mountain Basalt (Lower Jurassic) (Olsen, 1980a) - Dark-greenish-gray to black, fine-grained, dense, hard basalt composed mostly of calcic plagioclase and clinopyroxene. Locally contains small tubular gas-escape vesicles, some filled by zeolite minerals or calcite, and small to large vesicles lined with prehnite typically above base of flow contact. Unit consists of three major flows that are separated by a weathered zone, a bed of thin copper-sulfide-bearing, reddish-brown siltstone, or by volcanoclastic rock. Upper part of flow marked by olive-green hydrothermally altered horizon. Lower part of upper flow is locally planar-bedded, upper flow is massive to columnar jointed. Middle flow is massive to columnar jointed. Lower flow is generally massive with widely spaced curvilinear joints and is pillowed near the top. Individual flow contacts characterized by vesiculated zones as much as 8 ft thick. Reactivated lava tubes occur around pillow structures. Thickness of unit is about 715 ft.
- Jjp** Passaic Formation (Lower Jurassic and Upper Triassic) (Olsen, 1980a) - Interbedded sequence of reddish-brown, and less commonly maroon or purple, fine- to coarse-grained sandstone, siltstone, shaly siltstone, siltly mudstone, and mudstone. Reddish-brown sandstone and siltstone are thin- to medium-bedded, planar to cross-bedded, micaceous, and locally mudcracked and ripple cross-laminated. Root casts and load casts are common. Shaly siltstone, siltly mudstone, and mudstone are fine-grained, very thin to thin-bedded, planar to ripple cross-laminated, locally fissile, botryoidal, and contain evaporite minerals. They form rhythmically fining-upward sequences as much as 15 ft. thick. Several inches of unit have been thermally metamorphosed and locally mineralized with copper sulfides along contact with Orange Mountain Basalt (Jo). Unit is barely exposed in southwestern part of the map area, but regionally is as much as 1,450 ft. thick.

**EXPLANATION OF MAP SYMBOLS**

- Contact - Dashed where approximately located; dotted where concealed; queried where uncertain
- Faults - U, upthrown side; D, downthrown side; Ball shows direction of dip
- Dashed where approximately located
- Fault - High-angle fault of unknown movement
- Fold axes - Showing trace of axial surface, direction and dip of limbs
- Syncline - Showing trough
- Anticline - Showing crestline
- Plane features
- Strike and dip of inclined beds
- Other features
- Drill hole locations listed in Table 1
- Abandoned rock quarry
- Active rock quarry

Table 1. Well information from New Jersey Department of Environmental Protection well permit database.

Well Number	Well Permit Number	Depth (ft) below grade	Driller's Log	Remarks
1	25-23739	0-4	overburden	red shale
2	25-20475	0-370	overburden	red shale
3	25-24588	0-5	overburden	red shale
4	25-16110	0-10	shale	red shale
5	25-20041	0-12	overburden	red shale
6	25-22614	0-15	overburden	red shale
7	25-20227	0-12	overburden	red shale
8	25-20305	0-12	overburden	red shale
9	25-17119	0-195	trap rock/baked shale contact	trap rock
10	25-20307	0-13	overburden	trap rock
11	25-25559	0-8	overburden	trap rock
12	25-25075	0-40	broken rock	trap rock
13	25-23904	0-136	overburden	trap rock
14	90-00136	0-65.5	overburden	trap rock
15	25-25263	0-15	overburden	trap rock
16	25-25037	0-15	overburden	trap rock
17	25-06838	0-21	overburden	trap rock
18	25-19602	0-140	sandstone	trap rock
19	25-26235	0-30	overburden	trap rock
20	25-20337	0-30	overburden	trap rock
21	25-10108	0-56	clay	trap rock
22	25-00942	0-56	overburden	trap rock
23	25-00180	0-3	overburden	trap rock
24	25-17255	0-15	overburden	trap rock
25	25-13238	0-70	overburden	trap rock
26	25-10269	0-22	rotten rock	trap rock
27	25-10194	0-12	overburden	trap rock
28	25-19177	0-40	overburden	trap rock
29	25-10363	0-80	decomposed shale & rock	trap rock
30	25-00485	0-132	overburden	trap rock
31	25-10630	0-72	overburden	trap rock
32	25-03271	0-125	overburden	trap rock
33	90-00113	0-223	overburden	trap rock
34	25-05068	0-186	overburden	trap rock

**CORRELATION OF MAP UNITS**

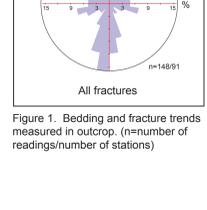
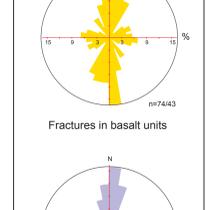
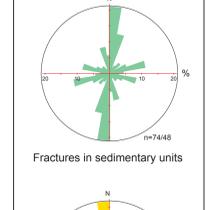
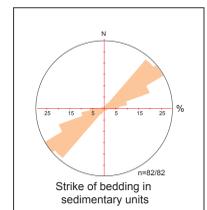
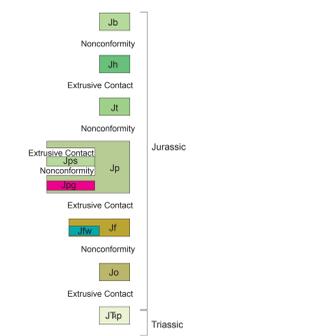


Figure 1. Bedding and fracture trends measured in outcrop. (n=number of readings/number of stations)

**Bedrock Geologic Map of the Chatham Quadrangle**  
Morris, Somerset and Union Counties, New Jersey  
By  
Donald H. Monteverde and Richard A. Volkert  
2005

