

ABSTRACT  
Bouguer gravity anomalies, residual gravity anomalies, and modeling of gravity profiles have been used to delineate the subsurface configuration of the main body of the Beermville intrusive complex and the enclosing Paleozoic sedimentary rocks. Surface geologic data were used to constrain the models and aeromagnetic maps were used to assist in interpretation of anomalies.

A gravity high associated with the intrusive extends southeast from Beermville. North-northwest-south-southwest trending gravity lows are associated with faulting within Proterozoic and Paleozoic rocks east of the study area.

The Beermville intrusive complex appears to be thin near the surface, and to broaden and elongate with depth. It extends to the west and probably plunges at a high angle to the southeast. Paleozoic sediments deepen and thicken to the northeast.

INTRODUCTION  
The Beermville intrusive complex in Sussex County, New Jersey consists of carbonaceous-silicic igneous rocks of Late Ordovician age. These are restricted in regional occurrence but vary widely in petrologic and intrusive character. Many previous studies have addressed the age, petrology, and morphology of the complex, but little has been done on its regional geology, particularly with its vertical dimension. The purpose of this study is to delineate the subsurface configuration of the main body of the intrusive complex using gravity data in conjunction with aeromagnetic data and surface geologic information.

An area of about 384 square miles (914 km<sup>2</sup>) was surveyed using the gravity method. Gravity observations were collected at 156 stations. These data were supplemented by gravity observations from 176 locations provided by Walter J. Spink (written communication, 1986).

GENERAL GEOLOGY  
The Beermville intrusive complex in the Kittatinny Valley section of the Valley and Ridge Physiographic Province. Paleozoic sedimentary rocks overlie Middle Proterozoic metamorphic and igneous rocks (Fig. 1, table 1). Surficial sediments are absent or only a few feet thick across most of the study area, thicker within an area extending north from Culvers Lake east to Onawissa. Large part these data to the late Wisconsin glacialation.

To the east, the New Jersey Highlands are underlain primarily by Proterozoic metamorphic and igneous rocks including gneiss, hornblende gneiss, schists, marbles and slates, and amphibolite. Similar areas are underlain by thickened Paleozoic sedimentary rocks unconformably overlying Proterozoic rocks. In both Kittatinny Valley and the Highlands, northeast-trending, moderate-to-high-angle faults cut both the Paleozoic and Proterozoic rocks. Detailed studies of the bedrock geology include Merchant and Treat (1954), Hague and others (1955), Spink (1967), Baker and Buddington (1970), Markewicz and Dalton (1977), Hernan and Morawski (1989) and Morawski (in press).

The Beermville complex intrudes carbonate rocks ranging in age from Proterozoic to Late Ordovician, but crops out generally within the Middle to Upper Ordovician Marlburg Formation. Radiometric dating of feldspar from the repetitive zircon at Beermville yields an age of about 435 ± 20 m.y. for the intrusion (Zarnham and others, 1987). Radcliffe (1981) estimated this age to 440 m.y.

The primary intrusive bodies at the surface are two nepheline syenite plutons, located along the contact between the Marlburg and Shawangunk Formations, and a rhyolite dike within the Marlburg. Other intrusive bodies include Phonolite, tragiolite and lamprophyre dikes, diatremes, and possibly alkali. Many (1970) reported a potassic schist for the intrusive complex which involved immiscibility and the fractional crystallization of a carbonate parent magma of basic composition. The complex is probably post-tectonic relative to the deformation of Ordovician rocks (Radcliffe, 1981). Ordovician deformation in the vicinity of the Beermville complex is characterized by open, upright folds within the Kittatinny Supergroup and Marlburg Formation (Hernan and Morawski, 1989). Subsequent Algonquin orogenesis complicates the structural setting.

GRAVITY FIELD METHODS  
A Lascaris & Rosenberg integral gravity method (model 020) capable of being read to the nearest 0.001 milligal (mGal) was used in this study. Secondary base stations were established at several locations in the area. These were referenced to a primary base station at Princeton University (Bonin and Woodard, 1957) with repeated visits.

Gravity observations by Spink (written communication, 1986) were taken using a Wallace Fleming gravimeter (model G133). The station elevations were determined using a Wallace and Tiernan altimeter referenced to benchmarks.

GRAVITY DATA REDUCTION  
Gravity observations at each station were corrected for instrumental drift, tide, latitude and elevation effects using standard formulas from Dobin (1976). The International Gravity Formula of 1930 (Dubin, 1976) was used to determine the theoretical gravity at sea level at a given latitude. A density of 2.67 g/cm<sup>3</sup> was used for Bouguer corrections.

Terrain corrections, which account for the deviation of topography from a horizontal surface, were calculated using a two-dimensional terrain correction computer program. The program calculates the vertical component of the gravity anomaly using the free-integral algorithm of Talwar and others (1959). Horizontal components of the gravity from the Bouguer slab effect (Dubin, 1976) results in an approximation of the terrain-correction value. Gravity terrain-correction values were calculated using 17 parallel profiles oriented north-northwest-south-southwest, with stations at 1640-foot (500-meter) intervals. Profiles were 1.9 miles (3 km) apart. The terrain-correction values were plotted and contoured to obtain a gravity terrain-correction map. Terrain corrections of several stations on one profile were determined using a Hammer graticule (Hammer, 1933) to verify the two-dimensional method. The correction factor difference between the two techniques ranged from 0.4 to -0.1 milligal. The terrain-correction factor for each gravity station was interpolated from the terrain-correction map and was added to the gravity values to obtain the Bouguer gravity anomaly. The terrain corrections ranged in value from 0.07 to 0.20 mGal.

The Bouguer gravity anomaly values (table 2) were then plotted and contoured to obtain a Bouguer gravity anomaly map (Fig. 1).

GRAVITY DATA ANALYSIS  
The Bouguer gravity values were analyzed using the double Fourier series analysis program of Jones (1968). This program is used for trend surface analysis of irregularly spaced data. A trend surface approximates the regional gravity. With the first harmonic, the analysis removed from the Bouguer gravity values at seawards of 53 miles (85 km) in the north-south direction and 27.5 miles (44 km) in the east-west direction. The residual gravity values thus obtained were then plotted and contoured to obtain the residual gravity map (Fig. 2).

A two-dimensional, nonlinear, least-squares gravity inversion program was used to model the Bouguer gravity data to profile form. This program incorporates the Marquadt procedure (Beck and Arnold, 1977) to calculate and optimize new parameters for each iteration of the model. A straight-line regional gravity gradient was also calculated. The gravity field of the initial model was calculated using the free-integral algorithm of Talwar and others (1959). Three profiles (Fig. 1) were modeled. The interpretation of gravity anomalies is not unique, hence requires some constraints in the modeling process. In this case the surface geology and the density contrasts of the outcropping geologic units were constrained. The depths to Middle Proterozoic gneiss and granitic basement and Paleozoic sedimentary units were extrapolated from the surface geologic structure. Thicknesses were from nearby measured sections. Densities were adopted from Dobin (1976). Tallied values (1976), and from rock sample measurements. These were contrasted with the assumed density of Middle Proterozoic basement of 2.67 g/cm<sup>3</sup>. Density contrasts (g/cm<sup>3</sup>) are:

Table with 2 columns: Unit and Density contrast (g/cm<sup>3</sup>). Rows include Quaternary glacial sediments, Devonian sedimentary rocks, Silurian quartzite/conglomerate, Ordovician slate and graywacke, Cambrian and Ordovician carbonates, Beermville intrusives and related rocks, Volcanic breccia and related rocks, and Mafic rocks.

RESULTS AND INTERPRETATION  
The Bouguer gravity anomaly map and the residual gravity map show a closed gravity high in the vicinity of the Beermville intrusive complex outcrop. An aeromagnetic high is also visible in the same area (Fig. 3). The gravity high results from the contrast between intrusive bodies and the lower density Paleozoic sedimentary rocks. The gradient of the gravity closure is gentle in the southeast and steep in the northwest, suggesting a southwesterly dip of the intrusive body. A second gravity high in the Hamonyville-Lewisburg area may be due to a part of the Beermville intrusive or to a change in Proterozoic rock density in the subsurface.

MAP SYMBOLS  
MAP SCALE  
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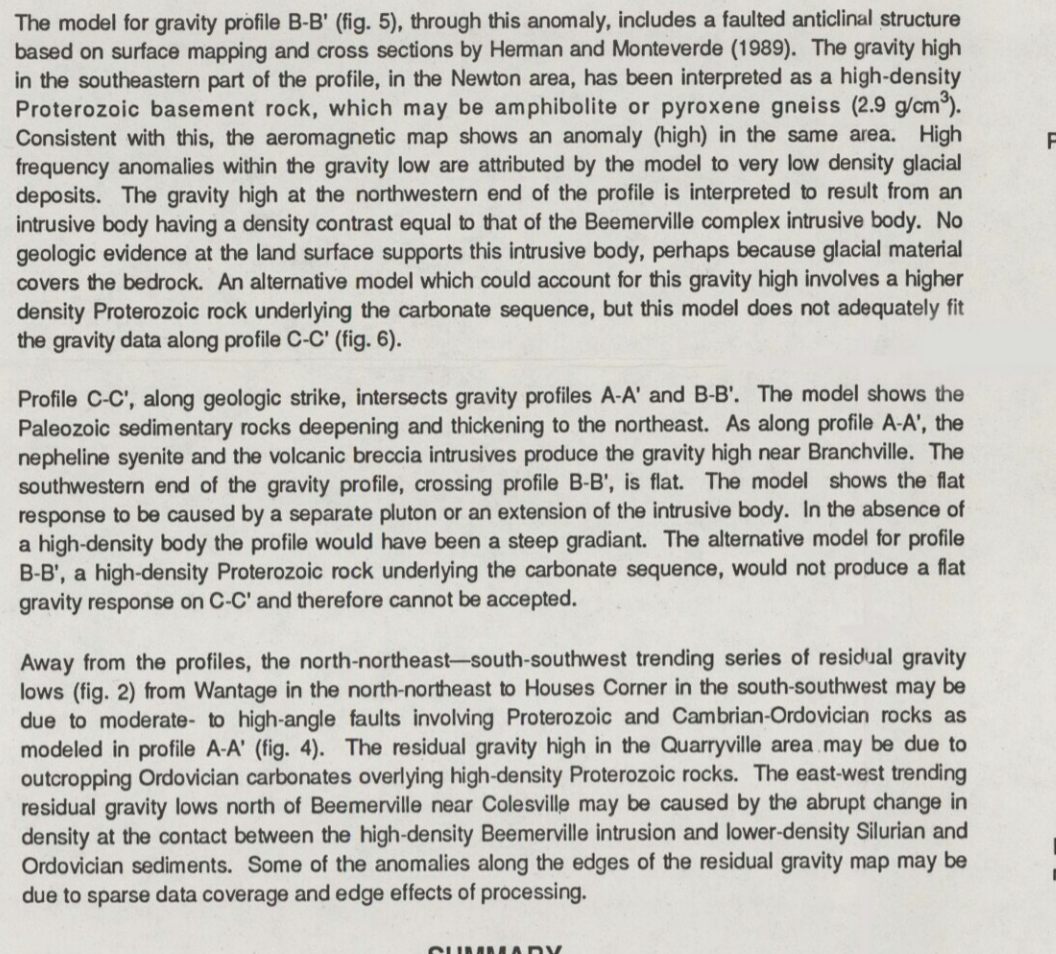


Table 1. Paleozoic lithostratigraphic units. Columns: System, Lithostratigraphic Unit, Average Thickness (ft), Description.

A closed gravity high is associated with the Beermville intrusive complex, resulting from the higher density of the intrusives as compared to the host Paleozoic sedimentary rocks. The Beermville intrusive complex appears to extend to the southwest. Two-dimensional modeling shows that the intrusives is elongate in shape and dips nearly vertically to the northeast. Paleozoic sedimentary rocks appear to deepen and thicken from south to north within the study area. A series of north-northwest-south-southwest trending gravity lows within the New Jersey Highlands result from downwarping of Paleozoic sedimentary rocks in areas of Proterozoic rock.

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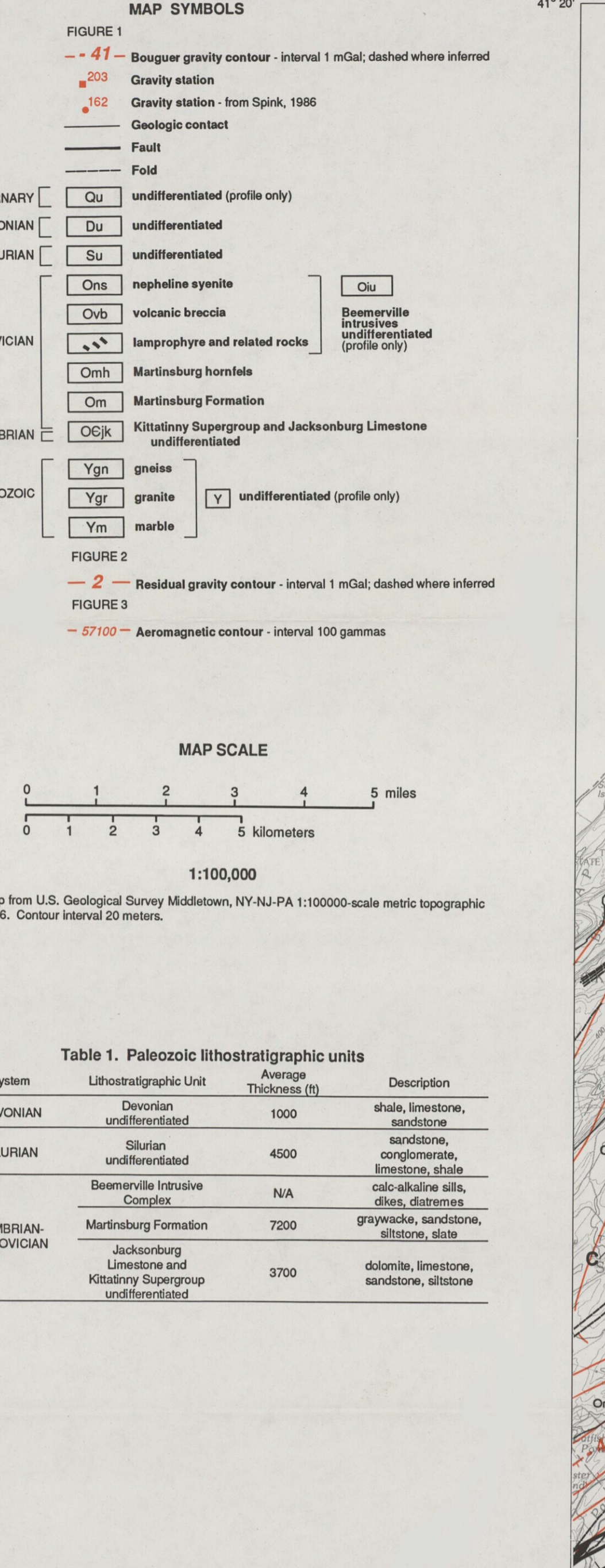


Figure 1. Bouguer gravity anomaly and geologic map.

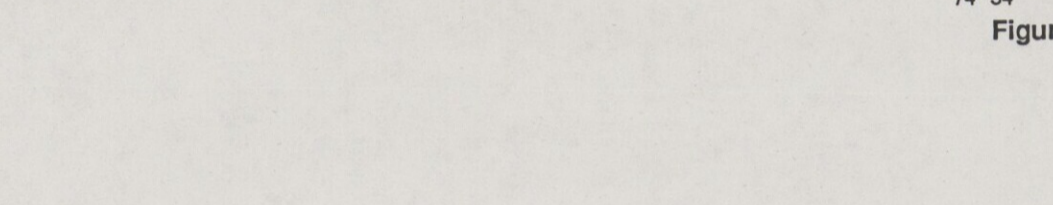


Figure 2. Residual gravity map.

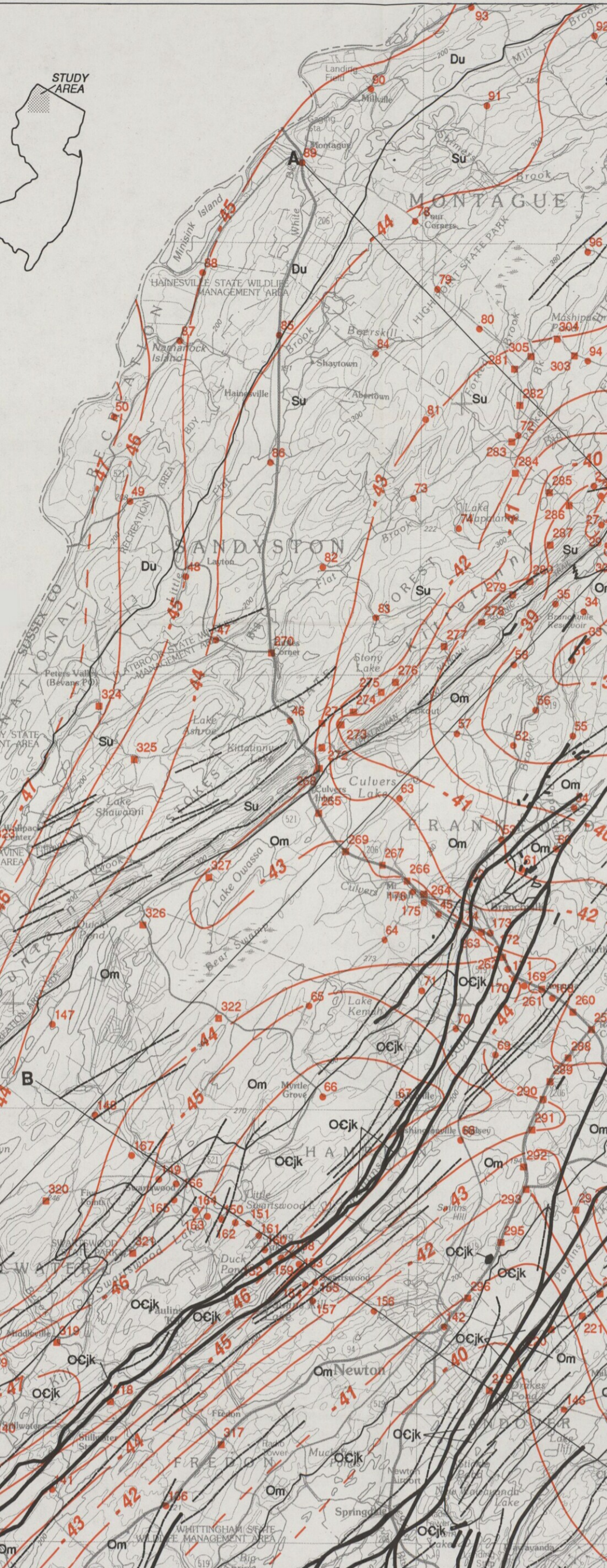


Figure 3. Aeromagnetic map.

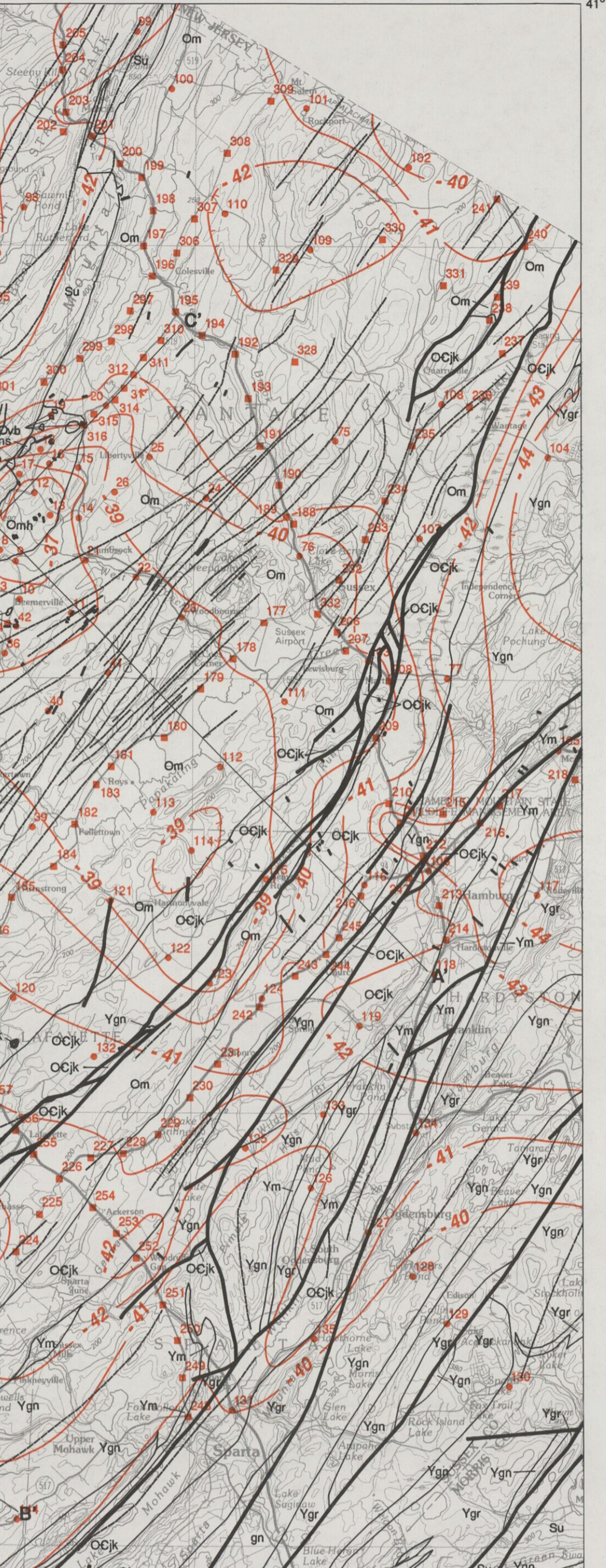


Figure 4. Bouguer gravity anomaly and model of profile A-A'. Line of section shown in Figure 1.

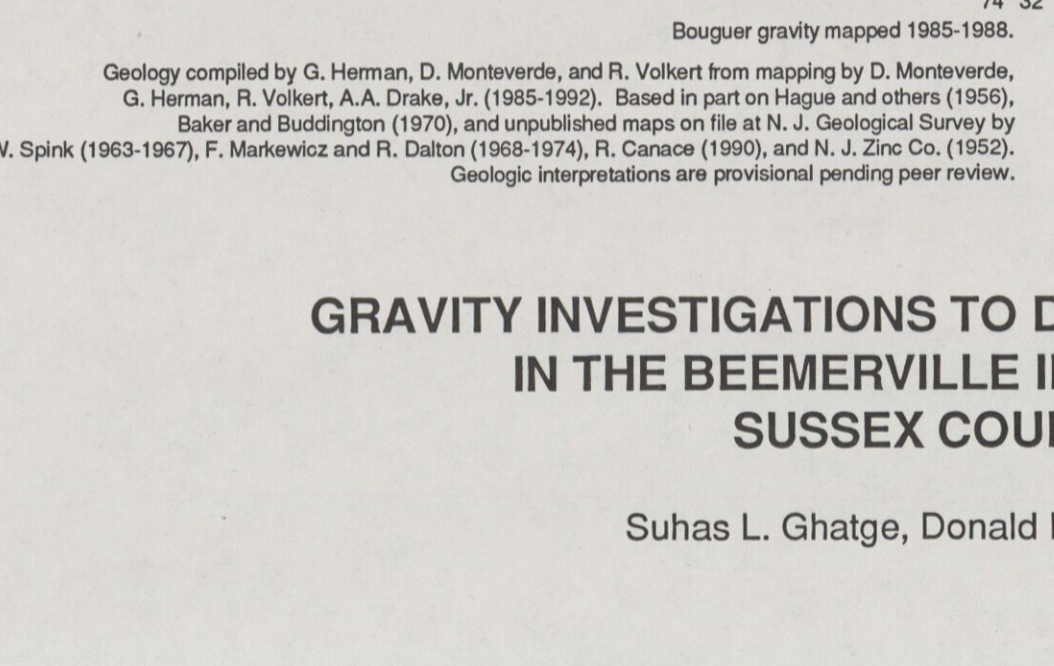


Figure 5. Bouguer gravity anomaly and model of profile B-B'. Line of section shown in Figure 1.

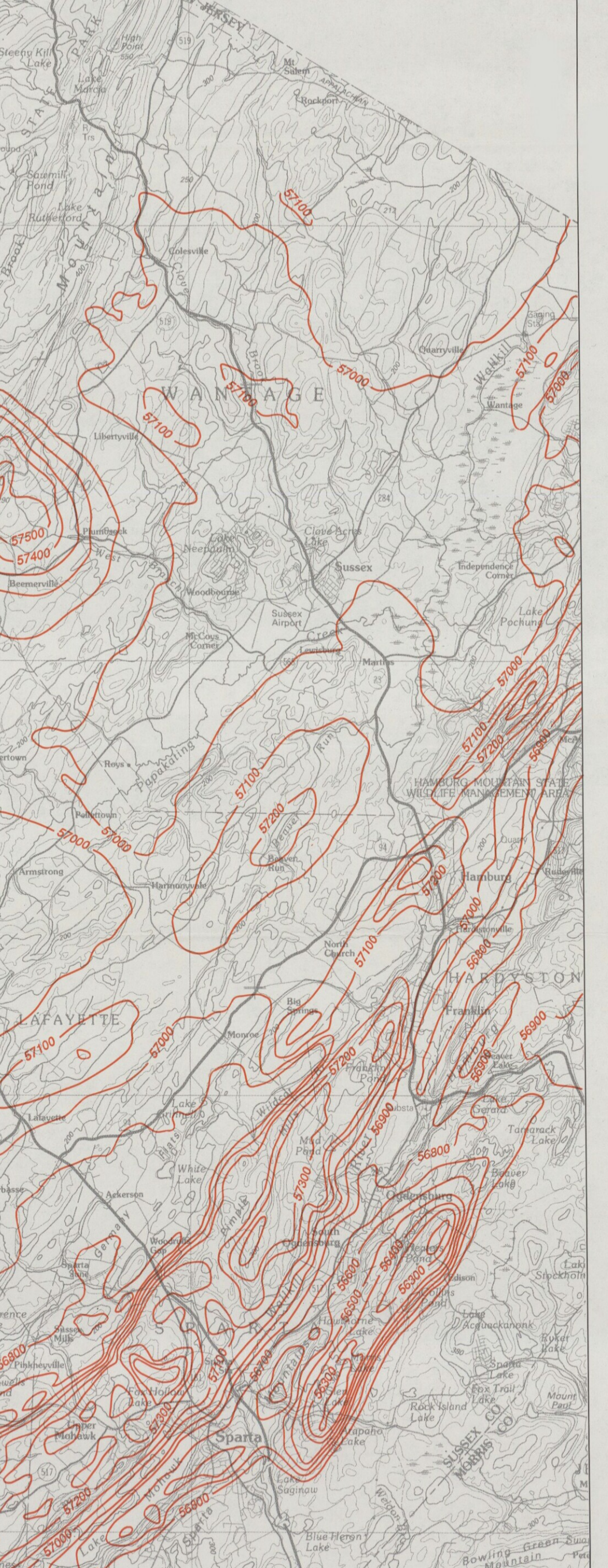


Figure 6. Bouguer gravity anomaly and model of profile C-C'. Line of section shown in Figure 1.

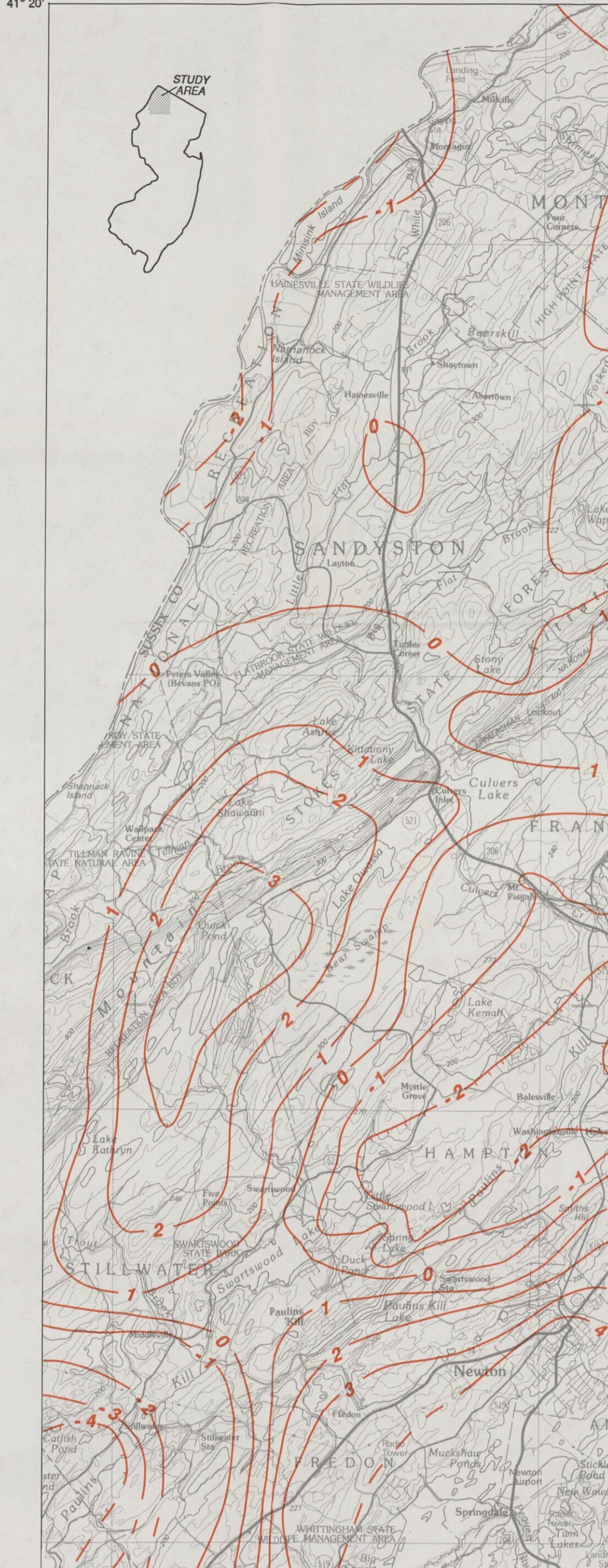


Figure 7. Bouguer gravity anomaly and geologic map.

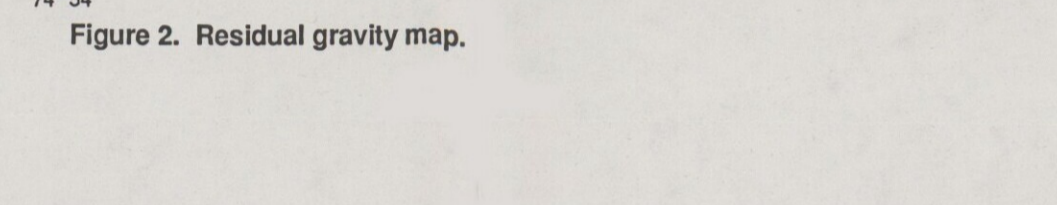


Figure 8. Residual gravity map.

GRAVITY INVESTIGATIONS TO DELINEATE SUBSURFACE GEOLOGY IN THE BEEMERVILLE INTRUSIVE COMPLEX AREA, SUSSEX COUNTY, NEW JERSEY

Suhans L. Ghatge, Donald L. Jagel, and Gregory C. Hernan 1992

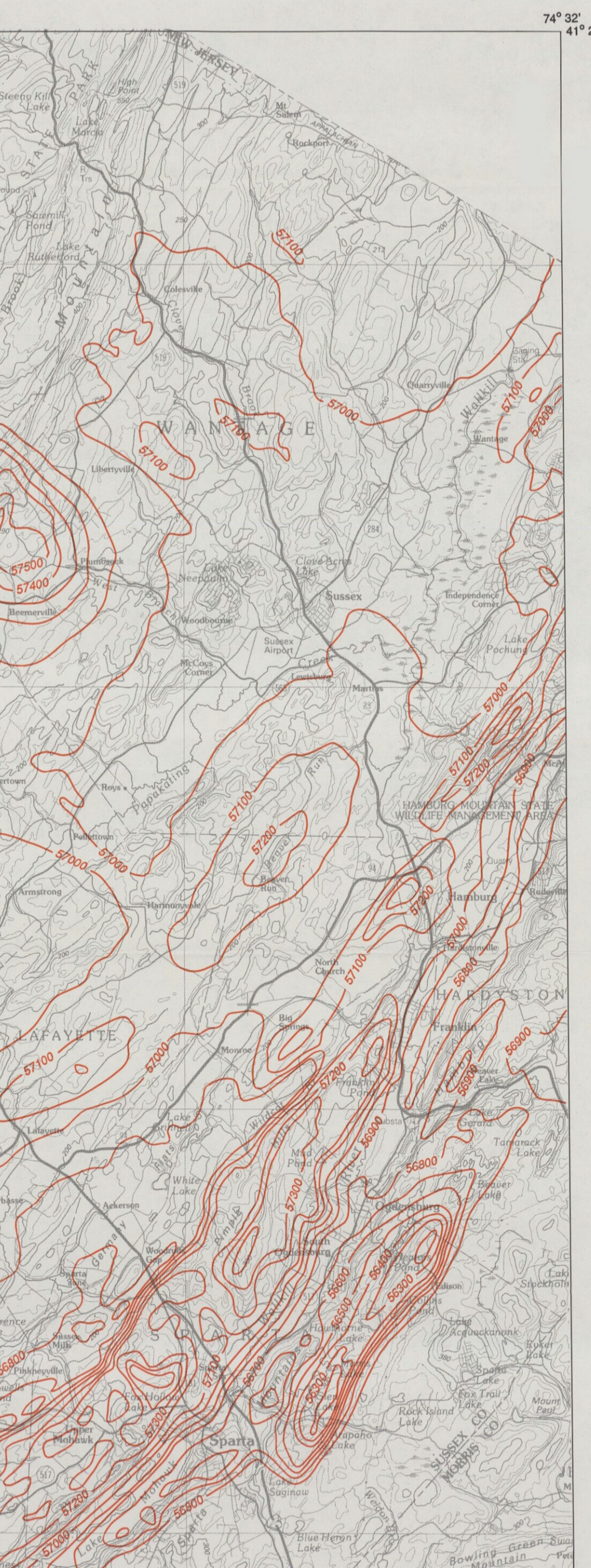


Figure 9. Bouguer gravity anomaly and geologic map.