25'

 Qt

S(O

UTH AMBO

NCT OI N) HETTPH *UTH JU OMN O M(*25' Kmo XO.

-
-
-
- clearly visible on 1:12,000-scale aerial photographs taken in 1979. Other basins, now destroyed by urbanization or obscured by thick vegetation, are shown by the base-map topography. These basins are likely of late Wisconsinan periglacial origin (Wolfe, 1953).

Trs

Trl | Lockatong Formation

Trlh Lockatong Formation, hornfels

GEOLOGY OF THE NEW BRUNSWICK QUADRANGLE, MIDDLESEX AND SOMERSET COUNTIES, NEW JERSEY

By:

Scott D. Stanford¹, Donald H. Monteverde¹, Richard A. Volkert¹, Peter J. Sugarman¹ And Gilbert J. Brenner²

1998

(for cross sections, see Sheet 2)

LOCATION MAP

GEOLOGY OF THE NEW BRUNSWICK QUADRANGLE MIDDLESEX AND SOMERSET COUNTIES, NEW JERSEY GEOLOGIC MAP SERIES OFM-23 SHEET 1OF 3

z (Feet)

 -1000

 -1000

 -2000

- 3000

 B°

ELEVATIO

DEPARTMENT OF ENVIRONMENTAL PROTECTION DIVISION OF SCIENCE AND RESEARCH NEW JERSEY GEOLOGICAL SURVEY

 $300 -$

 $200 -$

 10^c

Prepared in cooperation with the U.S. GEOLOGICAL SURVEY NATIONAL GEOLOGIC MAPPING PROGRAM

GEOLOGY OF THE NEW BRUNSWICK QUADRANGLE MIDDLESEX AND SOMERSET COUNTIES, NEW JERSEY GEOLOGIC MAP SERIES OFM-23 SHEET 2 OF 3

GEOLOGY OF THE NEW BRUNSWICK QUADRANGLE,

DEPARTMENT OF ENVIRONMENTAL PROTECTION DIVISION OF SCIENCE AND RESEARCH. NEW JERSEY GEOLOGICAL SURVEY

INTRODUCTION

Bedrock in the New Brunswick quadrangle includes shale and mudstone of Triassic age, intruded by diabase of Jurassic age. These rocks dip gently to the northwest, and crop out in the northwestern part of the quadrangle. They unconformably overlie schist and gneiss of Middle Proterozoic to Cambrian age, which occur in the subsurface in the southeastern part of the quadrangle. The bedrock formations are onlapped to the southeast by sand and clay of Cretaceous age. The Cretaceous sediments dip very gently to the southeast and crop out in the central and southeastern parts of the quadrangle. The bedrock and Cretaceous sediments are discontinuously overlain by surficial deposits of sand, gravel, mud, and peat of late Tertiary and Quaternary age.

These rocks and sediments are described in detail below. Table 1 provides drillers' logs of selected wells and test borings. These data show the thickness of surficial deposits and the underlying bedrock or Cretaceous formation. Figures 1 and 2 provide data on joints in the bedrock units.

DESCRIPTION OF MAP UNITS

Surficial Deposits

Surficial deposits are mapped where they are generally continuous and greater than 3 feet thick. Where no surficial deposits are mapped, discontinuous colluvium and alluvium, generally less than 3 feet thick, may overlie the mapped bedrock or Cretaceous formations. Bedrock units generally have a mantle of weathered material overlying unweathered rock. This material is not mapped separately from the bedrock formation. On the Passaic and Lockatong Formations, the weathered material is a reddish brown, reddish yellow, and gray clayey silt with variable amounts of angular rock fragments. It is generally less than 10 feet thick. On diabase, it is a reddish yellow sandy clay with many angular rock fragments and some to many rounded diabase boulders, also generally less than 10 feet thick. The Cretaceous deposits are oxidized and mineralogically altered but weathering has not significantly changed their grain size. Throughout the quadrangle, gravel and matrix material derived from the Pensauken Formation have been mixed into the upper several feet of the weathered material by soil processes.

- ARTIFICIAL FILL--Excavated sand, silt, clay, gravel, rock and man-made materials (bricks, cinders, ash, construction debris). In railroad and road embankments, dams, and made land. As much 60 feet thick. Many small areas of fill are not mapped. Extent of fill based on aerial photographs taken in 1979 and 1986.
- TRASH FILL-Trash and construction debris mixed and covered with clay, silt, and sand. As much as 100 feet thick.
- Qal ALLUVIUM--Sand, silt, pebble-to-cobble gravel, minor clay and peat. Contains variable amounts of organic matter. Matrix is reddish brown, yellowish brown, dark brown, light gray. The sand is predominantly quartz, shale fragments, and mudstone fragments; with variable but minor glauconite, mica, and feldspar. The gravel consists of shale, mudstone, quartz, quartzite, ironstone; and chert, diabase. and sandstone. As much as 15 feet thick (estimated). Alluvium along streams draining shale terrain has reddish brown fine sediment with much shale in both the sand and gravel fraction. Alluvium along streams draining Cretaceous sediments is brown to light gray and contains chiefly quartz and mica in the sand fraction. Older surficial deposits (units Tp, Qtu, and Qtl) contribute glauconite and feldspar to the sand fraction, and quartz, quartzite, ironstone, chert, and sandstone to the gravel fraction.
- Qm **ESTUARINE DEPOSITS**--Peat and organic-rich clay and silt, brown to dark gray; minor sand and shells. As much as 35 thick.

 \overline{Qs} SWAMP DEPOSITS--Peat and organic-rich clay and silt, minor sand, brown to dark gray. As much as 10 feet thick. Wood and peat from the base of swamp deposits in a basin near Fresh Ponds vielded radiocarbon dates of 9805±62 yrs B. P. (GX-17456), and 6910±120 yrs B. P. (I-16768) from depths of 5.5 and 3.5 feet respectively. Peat from the base of swamp deposits in a second basin half a mile to the south yielded a radiocarbon date of 3850±100 yrs B. P. (1-16769) from a depth of 4 feet. Peaty sand at a depth of about 6 feet at the bottom of organic sediments in the bog around Helmetta Pond (exact location unknown) yielded a date of 9649±60 yrs B. P. (QL-1082) (Watts, 1979). These dates indicate that swamp deposits have been gradually accumulating since at least the early Holocene.

 $|Qt|$ LOWER TERRACE DEPOSITS--Sand, silt, pebble gravel, and minor clay forming stream terraces with surfaces 5 to 20 feet above modern alluvial surfaces. Matrix is reddish brown, yellow, brownish yellow, and reddish yellow. Sand and gravel composition is similar to unit Qal, although organic matter is absent. The dependence of color and sand and gravel composition on source terrain is also similar to that of Qal. As much as 40 feet thick. In the Matchaponix and Manalapan Brook valleys, the lower terrace deposits are chiefly sand. Here, they are particularly rich in glauconite because they are derived, in part, from glauconitic Cretaceous sediments to the south and southeast of the quadrangle.

> Qtl is on grade with Qrt, which, in turn, is on grade with glaciofluvial deposits of late Wisconsinan age both upstream and downstream in the Raritan Valley. These relationships indicate that Qtl is, in part, of late Wisconsinan age. The Qtl deposits were assigned to the Cape May Formation by Lewis and Kummel (1910) and Salisbury and Knapp (1917). The name is not used here because the Cape May has been subdivided and redefined in its type area (Newell and others, 1989) and correlation of Qtl to the redefined Cape May deposits has not been established.

- **Qcsg** SAND-AND-GRAVEL COLLUVIUM--Sand, pebble gravel, minor silt and clay. Matrix is light gray, very pale brown, yellow and yellowish brown. Sand is chiefly quartz, with some mica and feldspar, and minor glauconite. Gravel is chiefly quartz, quartzite, ironstone; with minor chert, sandstone, and mudstone. As much as 20 feet thick. Forms aprons with surfaces that grade distally to elevations 5 to 20 feet above modern alluvial surfaces. Contemporaneous with unit Qtl.
- SHALE COLLUVIUM-Silt, clay, pebble gravel, minor sand. Matrix is reddish brown to brown. Sand is chiefly quartz, shale fragments, and feldspar. Gravel is shale, quartz, quartzite; with minor ironstone, chert, and sandstone. As much as 10 feet thick (estimated). Topographic and age relationships as in unit Ocsg.
- $Q\pi$ RARITAN TERRACE DEPOSIT-Sand, silt, pebble gravel, minor clay and cobble gravel. Matrix is gray, brown, and reddish brown. Sand is predominantly quartz, with some shale fragments and feldspar, and minor glauconite and mica. Gravel is predominantly quartz and quartzite; with some red and gray mudstone and shale; and minor chert, gneiss, and sandstone. As much as 20 feet thick.

Unit Qrt is on grade with late Wisconsinan glaciofluvial deposits both upstream and downstream in the Raritan Valley and so is, in part, of late Wisconsinan age. It likely includes some glacially derived sediment from bedrock to the north and east of the Raritan basin, in addition to nonglacially derived sediment from bedrock, Coastal Plain formations, and surficial deposits within the basin. The deposit was assigned to the Cape May Formation by Darton and others (1908) and Lewis and Kummel (1910). The name is not used here because the Cape May has been subdivided and redefined in its type area (Newell and others, 1989) and correlation of Qrt to the redefined Cape May deposits has not been established.

- | Qtu | UPPER TERRACE DEPOSITS--Sand, yellow to reddish yellow; pebble gravel, minor cobble gravel. On terrace remnants with bases that are 20 to 40 feet above modern alluvial and salt-marsh surfaces. Sand and gravel composition similar to that of unit Tp. As much 20 feet thick. Topographic position indicates that unit Otu is older than units Qtl and Qrt but younger than unit Tp, and so may range in age from Early to Middle Pleistocene.
- **Tp PENSAUKEN FORMATION-Sand, reddish yellow to yellow;** pebble gravel; minor cobble gravel. Sand is predominantly quartz; with some feldspar; and minor glauconite, mica, and red shale. Gravel is predominantly quartz, quartzite, and ironstone; with chert, red to gray mudstone, and sandstone; and minor gneiss, schist, and diabase. As much as 80 feet thick. Cobble-gravel channel deposits are common in the basal few feet of the deposit, where they are commonly iron-cemented and contain abundant clasts of sandstone and mudstone, and scattered clasts of gneiss, schist, and diabase. These clasts generally have thick weathering rinds or are fully decomposed. Iron-petrified wood also occurs in the basal few feet of the deposit in places. Tabular, planar cross-bedded sand with minor pebble gravel dominates the deposit above the basal gravel. The pebble gravel is chiefly quartz, quartzite, and ironstone with some chert and minor mudstone. A red-colored soil is developed in the Pensauken on flat surfaces above about 140 feet in elevation in the Lawrence-Brook-Manalapan-Brook divide area. These upland flats may be remnants of the original aggradation surface. Other surfaces on the Pensauken are not flat and do not exhibit a red soil, indicating that they were formed by stream dissection and slope erosion during the Pleistocene.

Salisbury and Knapp (1917) defined and mapped the Pensauken Formation. Owens and Minard (1979) reassigned the Pensauken deposits north of Trenton to the Bridgeton Formation (a higher fluvial sand and gravel in southern New Jersey), based on projection of the elevations of the deposits from their type areas in southern New Jersey. This usage was followed by Martino (1981) and Stanford (1993, 1995). However, the deposits north of Trenton are continuous in both extent and elevation with those at the Pensauken type locality, so the original nomenclature is used here. The age of the Pensauken is not firmly established. Berry and Hawkins (1935) describe plant fossils in the Pensauken from the Ireland Brook valley near Beth Abraham Cemetery that they consider to be of early Pleistocene age. Owens and Minard (1979) assign a late Miocene age based on correlation to units in the Delmarva Peninsula. Pollen from a black clay bed within the Pensauken near Plainsboro, New Jersey, about 12 miles southwest of Old Bridge, includes cool-temperate to cold-temperate species and a few pre-Pleistocene species. This assemblage suggests a Pliocene age (Gilbert J. Brenner, written communication, December 1991). This age is also consistent with the geomorphic and stratigraphic relation of the Pensauken to late Pliocene or early Pleistocene till and to middle to late Miocene marine and fluvial deposits (Stanford, 1993).

Coastal Plain Formations

[Kmg] MAGOTHY FORMATION (Darton, 1893)-Quartz sand, white to yellow, micaceous, commonly interbedded with thin to thick, gray carbonaceous clay and silt. In places the clay is oxidized to white. red, and pink. Locally, the clay contains pyrite and lignite. Sand is typically cross-stratified, although laminated beds are common. In the Sayreville-South Amboy area, just east of the New Brunswick quadrangle, extensive clay-pit exposures permitted the naming and mapping of clay beds within this sequence, which was formerly included in the Raritan Formation (Cook, 1878; Ries and others 1904). Later, intervening sands in the same area were named as part of an aquifer investigation (Barksdale and others, 1943). Berry (1905) correlated the uppermost beds of the Raritan with the Magothy Formation of Maryland, based on their fossil flora. Owens and Sohl (1969), Wolfe and Pakiser (1971), Owens and others (1977), and Christopher (1979) provide more detailed biostratigraphic control for these units, and redefined additional beds of the former Raritan as Magothy. In the South Amboy area the Magothy now includes, from oldest to youngest, the following informal members: South Amboy Fire Clay, Old Bridge Sand, Amboy Stoneware Clay, Morgan beds, and Cliffwood beds (Owens and others, 1977). These members may extend into the New Brunswick quadrangle but exposure and subsurface data are not sufficient to map them. The Magothy is Turonian-Coniacian (?) to Santonian in age (Christopher, 1979, 1982). It is as much as 180 feet thick in this quadrangle.

RARITAN FORMATION (Cook, 1868)--Includes two informal members in this quadrangle: the Farrington Sand and the Woodbridge Clay. Another member, the "Raritan fire and potter's clay" of Cook (1878) and Ries and others (1904), underlies the Farrington Sand and was mapped by Cook (1878) and Ries and others (1904) in the northeastern part of the quadrangle. This unit includes a lower clay (the "potter's clay") which is predominantly a red, white, and gray clay derived from weathering of shale and mudstone of the Passaic and Lockatong Formations, and is included in those formations on this map; and an upper, discontinuous, gray sandy clay (the "fire clay") which is near the base of the Farrington Sand and is included in that member on this map.

[Krw] WOODBRIDGE CLAY--Clay and silt, dark gray, massive, with mica, lignite (typically fine grained), and pyrite. Locally interlaminated with white to yellow quartz sand. Small (less than 1) foot thick) lense-like masses and slabs of gray to brown siderite are common. The Woodbridge Clay is as much as 120 feet thick in this quadrangle.

> The Woodbridge has been assigned to the Complexipollis-Atlantopollis pollen assemblage zone of Cenomanian age (Christopher, 1979), or Pollen Zone IV. It has also been assigned to the upper Cenomanian based on the ammonites Metoicoceras bergquisti and Metengonoceras sp. (Cobban and Kennedy, 1990). Samples in this quadrangle (locations plotted on map) yielded Zone IV pollen and cysts of the dinoflagellates Cyclonephelium distinctium, Hystrichospheridium recurvatum, and Cleistosphaeridium sp.

 Krf FARRINGTON SAND--Quartz sand, white, yellow, and pink, micaceous, commonly interbedded with thin beds of angular granule gravel and minor, thin to thick gray clay and silt beds. It rests unconformably on weathered rocks of Mesozoic age, and is as much as 100 feet thick in this quadrangle. Lower parts of the Farrington may be time-equivalent to the Potomac Formation.

> The Farrington is Cenomanian in age (Christopher, 1979). It has been assigned to the upper part of Pollen Zone III by Christopher (1979). Samples of the Farrington from the adjacent South Amboy and Perth Amboy quadrangles yielded Zone IV pollen assemblages (written communication, Leslie A. Sirkin, Adelphi University, 1989). One sample collected within the quadrangle near South River from laminated gray silt and white to yellow fine sand in the transitional contact zone between the Woodbridge and the Farrington had species generally associated with older palynomorphs than typically seen in the Woodbridge Clay, and is assigned to Zone III-IV.

> > **Bedrock Formations**

Jd DIABASE (Lower Jurassic)--Fine-grained to aphanitic dikes (?) and sills and medium-grained, discordant, sheet-like intrusion of darkgray to dark greenish-gray, sub-ophitic diabase. Massive-textured. hard, and sparsely fractured. Composed dominantly of plagioclase, clinopyroxene, and opaque minerals. Contacts are typically finegrained, display chilled, sharp margins and may be vesicular adjacent to enclosing sedimentary rock. Exposed in map area in sills along the Raritan River east of Highland Park, in a dike approximately 1.9 miles south-southeast of Voorhees, and in the Rocky Hill diabase sheet along Oakeys Brook immediately north of Davidsons Millpond County Park. This sheet may be the southern extension of the Palisades sill. Beneath Cretaceous deposits it is weathered to a white, gray, or olive clay as much as 20 feet thick. The thickness of the Rocky Hill diabase in the quadrangle, known mainly from drill-hole data, is approximately 1,325 feet.

LOCKATONG FORMATION (Upper Triassic) (Kummel, 1897)--Cyclically deposited sequences of mainly gray to greenish-gray, and in upper part of unit, locally reddish-brown siltstone to silty argillite and dark-gray to black shale and mudstone. Siltstone is medium- to

ripple cross-laminations and locally abundant pyrite. Shale and mudstone are very thin-bedded to thinly-laminated, platy, locally containing desiccation features. Thermally altered to dark-gray to black hornfels (Trlh) where intruded by diabase. Thickness of hornfels directly related to thickness of intruded diabase. Lower contact gradational into Stockton Formation and placed at base of lowest continuous black siltstone bed (Olsen, 1980). Beneath Cretaceous deposits the Lockatong Formation is weathered to a gray, white, or olive clay as much as 20 feet thick. This material was mapped as the "Raritan fire and potter's clay" member of the Raritan Formation by Cook (1878) and Ries and Kummel (1904) but is included here with the Lockatong Formation. Maximum thickness of unit regionally is about 2,200 feet (Parker and Houghton, 1990). Approximately 1,800 feet thick in this quadrangle, including about 160 feet of hornfels. Best exposed southeast of Farrington Lake.

fine-grained, thin-bedded, planar to cross-bedded with mud cracks,

Trs STOCKTON FORMATION (Upper Triassic) (Kummel, 1897)--Regionally unit is interbedded sequence of gray, grayish-brown, or reddish-brown, medium- to fine-grained, thin- to thick-bedded, planar to cross-bedded arkosic conglomerate, sandstone, and reddish-brown shaly siltstone to mudstone. Lower contact is an erosional unconformity. Conglomerate and sandstone units are deeply weathered; siltstone and mudstone are generally less weathered. Maximum thickness of unit regionally may be as much as 3,000 feet (Parker and Houghton, 1990). Interpreted to be approximately 2,550 feet thick in this quadrangle. Not exposed in map area but known from water well records and drill-hole data.

CZYu UNDIFFERENTIATED PRE-MESOZOIC BASEMENT (Cambrian to Middle Proterozoic (?)) (Volkert and others, in press) -Gray, medium- to coarse-grained schist and gneiss dominantly composed of quartz, feldspar, and mica (biotite and/or muscovite), deeply weathered; may be associated with green micaceous saprolite. Not exposed in map area but known from water well records and drill-hole data.

BEDROCK JOINTS

The distribution and orientation of joints in the bedrock units are shown on the "Joint Orientation Map" on sheet 1 and on figures 1 and 2. Joints are fractures that show no visible movement parallel to the surface of the fracture. Bedding partings are not considered. All joints were measured on outcrops. Strike and dip data were grouped by structural block. Structural blocks were delineated by identifying lithologic units that have undergone a similar deformational history. Because the lithology of the sedimentary rocks is similar, the structural blocks are delineated mainly by the major faults rather than by formation contacts. Diabase is classed separately because it is mechanically distinct from the sedimentary rocks. Joint orientations were analyzed by plotting rose diagrams of all joints within each structural block according to dip azimuth in 10-degree sectors (figure 1). The percentage of data in each sector as compared to the total domain database was then calculated using Field Data Management (FMS) software (G. C. Herman, Field Data Management System v. 2.1; an upgrade of Kaeding and Herman, 1988, on file at the N. J. Geological Survey). Percentage values were rounded off to whole numbers. Sectors without readings are omitted. Sectors showing 0 percent contain data but comprise less than 1 percent of the total. A module of the FMS software creates files for ARC/INFO Geographic Information Systems (GIS) programs, which assign a statistical weight factor according to the sector percentage values. The GIS programs recalculate the dip azimuth values back to strike and dip, then plot user-defined structural symbols using the weighting factor at specified locations on maps

 $160 p.$

Berry, E. W., 1905, The flora of the Cliffwood clays: N. J. Geological Survey Annual Report for 1905, p. 135-174. Berry, E. W, and Hawkins, A. C., 1935, Flora of the Pensauken Formation in New Jersey: Geological Society of America Bulletin, v. 46, p. 245-252.

Christopher, R. A., 1979, Normapolles and triporate pollen assemblages from the Raritan and Magothy Formations (Upper Cretaceous) of New Jersey: Palynology, v. 3, p. 73-121. Christopher, R. A., 1982, The occurrence of the Complexiopollis- Atlantopollis Zone (palynomorphs) in the Eagle Ford Group (Upper Cretaceous) of Texas: Journal of Paleontology, v. 56, p. 525-541.

Cobban, W. A., and Kennedy, W. J., 1990, Upper Cenomanian ammonites from the Woodbridge Clay member of the Raritan Formation in New Jersey: Journal of Paleontology, v. 64, p. 845-846. Cook, G. H., 1868, Geology of New Jersey: Geological Survey of New Jersey,

Cook, G. H., 1878, Report on the clay deposits of Woodbridge, South Amboy, and other places: Geological Survey of New Jersey, Trenton, N. J., 380 p. Darton, N. H., 1893, The Magothy formation of northeastern Maryland: American Journal of Science, 3d series, v. 45, p. 407-419.

Darton, N. H., Bayley, W. S., Salisbury, R. D., and Kummel, H. B., 1908, Description of the Passaic quadrangle: U. S. Geological Survey Geologic Atlas, Folio 157, 27 p.

Gronberg, J. M., Birkelo, B. A., and Pucci, A. A., 1989, Selected borehole geophysical logs and drillers' logs, northern Coastal Plain of New Jersey: U.S. Geological Survey Open-File Report 87-243, 133 p. Herman, G. C., French, M. A., and Monteverde, D. H., 1993, Automated mesostructural analysis using GIS, beta test: Paleozoic structures from the New Jersey Great Valley region: Geological Society of America Abstracts with Programs, v. 25, no. 2, p. 23.

College, $p. E1-E36$. Kaeding, Margaret, and Herman, G. C., 1988, Field data management system (FMS): a computer software program for organization and analysis of geologic data: N. J. Geological Survey Technical Memorandum 88-4, 48 p. Kummel, H.B., 1897, The Newark System, report of progress: New Jersey

states, p. 93-105.

Martino, R. L., 1981, The sedimentology of the late Tertiary Bridgeton and Pensauken Formations in southern New Jersey: unpublished Ph. D. thesis, Rutgers University, New Brunswick, N. J., 299 p.

 $p. 25-51.$

 $47 p.$ University Press, p. 235-278.

Parker, R. A., and Houghton, H. F., 1990, Bedrock geologic map of the Monmouth Junction quadrangle, Somerset and Mercer Counties, New Jersey: U.S. Geological Survey Open-File Report 90-219, scale 1:24,000.

Ries, Heinrich, Kummel, H. B., and Knapp, G. N., 1904, The clays and clay industry of New Jersey: N. J. Geological Survey Final Report of the State Geologist, v. $6, 548$ p.

Salisbury, R. D., and Knapp, G. N., 1917, The Quaternary formations of southern New Jersey: N. J. Geological Survey Final Report of the State Geologist, v. 8, 218 p.

Sandberg, S. K., Hall, D. W., Gronberg, J. M., Groenewold, J. C., and Pasicznyk, D. L., 1996, Geophysical investigation of the Potomac-Raritan-Magothy aquifer system and underlying bedrock in parts of Middlesex and Mercer counties, New Jersey: N. J. Geological Survey Report 37, 33 p.

Spayd, S. E., 1985, Movement of volatile organics through a fractured rock

aquifer: Ground Water, v. 23, p. 496-502. Stanford, S. D., 1993, Late Cenozoic surficial deposits and valley evolution of

unglaciated northern New Jersey: Geomorphology, v. 7, p. 267-288. Stanford, S. D., 1995, Surficial geology of the South Amboy quadrangle, Middlesex and Monmouth counties, New Jersey: N. J. Geological Survey Open File Map 18, scale 1:24,000.

Van Houten, F.B., 1969, Late Triassic Newark Group, north central New Jersey and adjacent Pennsylvania and New York, in Subitzky, Seymour, ed., Geology of selected areas in New Jersey and eastern Pennsylvania and guidebook of excursions: New Brunswick, New Jersey, Rutgers University Press, p. 314-347.

Vecchioli, John, 1967, Directional hydraulic behavior of a fractured-shale aquifer in New Jersey: Proceedings of the International Symposium on Hydrology of Fractured Rocks, International Association of Scientific Hydrology Publication 73, v. 1, p. 318-326.

Vecchioli, John, Carswell, L. D., and Kasabach, H. F., 1969, Occurrence and movement of ground water in the Brunswick shale at a site near Trenton, New Jersey: U. S. Geological Survey Professional Paper 650-B, p. 154-157. Volkert, R. A., Drake, A. A., Jr., and Sugarman, P. J., in press, Geology, geochemistry and tectonostratigraphic relations of the crystalline basement beneath the coastal plain of New Jersey and contiguous areas, in, Drake, A. A., Jr., ed., Geologic studies in New Jersey and eastern Pennsylvania: U.S. Geological Survey Professional Paper, chapter A.

Wolfe, J. A., and Pakiser, H. M., 1971, Stratigraphic interpretation of some

Wolfe, P. E., 1953, Periglacial freeze-thaw basins in New Jersey: Journal of Geology, v. 61, p. 131-141. Woodward, H. P., 1944, Copper mines and mining in New Jersey: Department of Conservation and Development State of New Jersey, Geologic Series **Bulletin** 57, 156 p.

Prepared in cooperation with the **U. S. GEOLOGICAL SURVEY** NATIONAL GEOLOGIC MAPPING PROGRAM

(Herman and others, 1993; M. A. French and G. C. Herman, MesoPlot v. 1.0, unpublished program on file at the N. J. Geological Survey). The longer the length of the symbol the more dominant the joint orientation. All plotted data on the map show strike trend. The bottom of the flag symbol marks the location of the outcrop. Dip values are divided into 30-degree sectors. Dips of 1 to 29 degrees are marked by open flags; half-filled flags mark dips of 30 to 59 degrees; filled flags marks dips of 60 to 89 degrees. Flag placement shows direction of dip. Some joints with identical strike trends and opposite dip direction will have different symbol lengths because the statistics were calculated on dip azimuth.

Joint data are important to fully characterize the water-bearing properties of the bedrock. Because the bedrock generally lacks primary porosity, joints and other fractures control the occurrence and movement of ground water. Models invoke either bedding-parallel fractures (Vecchioli and others, 1969; Michalski, 1990) or vertical fractures (Vecchioli, 1967; Spayd, 1985) as the dominant control on ground-water movement. Hybrid models suggest that either flow cannot be defined by bedding fractures or vertical joints alone but that both must be considered (Houghton, 1990; Lewis, 1992). All models show that the bedrock aquifers are anisotropic and ground-water flow is generally parallel rather than perpendicular to bedding strike direction. The same models characterize dominant joint trends as either parallel to the strike of bedding or perpendicular to it. The data presented here suggest that the dominant joint trends more closely parallel the strike of the major faults rather than strike of bedding (figure 2).

REFERENCES Barksdale, H. C., Johnson, M. E., Schaefer, E. J., Baker, R. C., and

DeBuchananne, G. D., 1943, The ground-water supplies of Middlesex County, New Jersey: N. J. State Water Supply Policy Commission Special Report 8,

Trenton, N. J., $900 p$.

Houghton, H. F., 1990, Hydrogeology of the early Mesozoic rocks of the Newark Basin, New Jersey: in, Kroll, R. L., and Brown, J. O., eds., Aspects of groundwater in New Jersey, field guide and proceedings of the seventh annual meeting of the Geological Association of New Jersey: Union, New Jersey, Kean

Geological Survey Annual Report of the State Geologist, 1896, p. 25-88. Lewis, J. C., 1992, Effect of anisotropy on ground-water discharge to streams in fractured Mesozoic-basin rocks: American Water Resources Association, regional aquifers of the United States, aquifers of the southern and eastern

Lewis, J. V., and Kummel, H. B., 1910, Geologic map of New Jersey: N. J. Geological Survey Atlas Sheet 40, scale 1:250,000. Revised by H. B. Kummel, 1931 and M. E. Johnson, 1950.

Michalski, Andrew, 1990, Hydrogeology of the Brunswick (Passaic) Formation and implications for ground water monitoring practice: Ground Water Monitoring Review, v. 10, no. 4, p. 134-143.

Newell, W. L., Wyckoff, J. S., Owens, J. P., and Farnsworth, John, 1989, Cenozoic geology and geomorphology of southern New Jersey Coastal Plain: U. S. Geological Survey Open-File Report 89-0159, 51 p.

Olsen, P.E., 1980, The latest Triassic and early Jurassic formations of the Newark Basin (eastern North America Newark Supergroup): stratigraphy, structure and correlation: New Jersey Academy of Science Bulletin, v. 25, no. 2,

Owens, J. P., and Minard, J. P., 1979, Upper Cenozoic sediments of the lower Delaware valley and northern Delmarva Peninsula, New Jersey, Pennsylvania, Delaware, and Maryland: U. S. Geological Survey Professional Paper 1067D,

Owens, J. P., and Sohl, N. F., 1969, Shelf and deltaic paleoenvironments in the Cretaceous-Tertiary formations of the New Jersey Coastal Plain, in Subitzky, Seymour, ed., Geology of selected areas in New Jersey and eastern Pennsylvania and guidebook of excursions: New Brunswick, N. J., Rutgers

Owens, J. P., Sohl, N. F., and Minard, J. P., 1977, A field guide to Cretaceous and lower Tertiary beds of the Raritan and Salisbury embayments, New Jersey, Delaware, and Maryland: Guidebook for the Annual American Association of Petroleum Geologists/Society of Economic Paleontologists and Mineralogists Convention, Washington, D. C., June 12-16, 1977, p. 58-69.

Watts, W. A., 1979, Late Quaternary vegetation of central Appalachia and the New Jersey Coastal Plain: Ecological Monographs, v. 49, p. 427-469.

Cretaceous microfossil floras of the Middle Atlantic States: U. S. Geological Survey Professional Paper 750-B, p. B35-47.

of clay (Kmg)

Fault trends

 $81-107$ gray coarse sand (Krf) 107-125 multicolored solid clay (weathered Trl)

of the unit (Sandberg and others, 1996). No samples were examined except as noted