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 \sim >200 feet Figure 2.--Thickness of surficial material, including weathered rock, in the Mendham quadrangle. Abbreviations for bedrock lithologies are: CB=carbonate rock, GR=unfoliated, massive granite, MG=foliated, micaceous gneiss (with marble lenses in the Whippany River valley), GN=foliated, layered gneiss, CS=conglomerate and shale, BA=basalt. Bedrock contacts and lithologies based on Volkert (2008). Thick surficial materials include Illinoian and late Wisconsinan glacial deposits in the Lamington, Mill Brook, Den

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Figure 1.--Glacial and geomorphic features of the Mendham quadrangle. Abbreviations on Illinoian glacial-lake shorelines are: DP=lake in valley south of Dalrymple Pond; JB1, JB2, JB3=high, intermediate, and low stage, respectively, of lake in the Jackson and Wallace Brook valleys; SH1, SH2, SH3=high, intermediate, and low stage, respectively, of glacial Lake Shongum. Abbreviations on late Wisconsinan glacial-lake shorelines are: SC=glacial Lake Succasunna, DO=glacial Lake Dover, MH=Moggy Hollow stage of glacial Lake Passaic. Abbreviations on Illinoian glacier margins are: M1=maximum extent of Illinoian ice, lake impounded in valley south of Dalrymple Pond and high stage of Lake Shongum impounded in Den Brook valley; M2=last ice margin before Lake Shongum, high stage, and high stage of lake in Jackson Brook valley lower to their intermediate stages; M3=last ice margin before intermediate stages of Lake Shongum and lake in Jackson Brook valley lower to their low stages. Topographic image from N. J. Highlands Water Protection and Planning Council Geographic

Brook, and Whippany valleys, colluvium at the base of escarpments and ridges, and saprolite developed on micaceous and foliated gneiss. Granite generally underlies the highest ridges and has the thinnest mantle of

weathered material.

Qaf Qal Qst $\overline{\mathsf{Qsl}}~\Big|$ Qdol $\Big|$ ^{Qpml} **Qcs Qcg Qcal Qwg Qwb Qwcb Qws Qf Qis weathered-rock material hillslope deposits glacial, fluvial, wetland, and manmade deposits hillslope erosion erosion and valley incision Holocene late Pleistocene (late Wisconsinan) late Pleistocene (early and middle Wisconsinan) middle Pleistocene Qso aft Qr Qs**

INTRODUCTION

The Mendham quadrangle is located in the Highlands province in north-central New Jersey. Surficial deposits in the quadrangle include glacial, stream, hillslope, wetland, and man-made sediment and weathered-rock material. Glacial sediments were laid down during three glaciations: the late Wisconsinan glaciation (about 25,000 years ago), an intermediate glaciation of pre-Wisconsinan age (most likely the late Illinoian glaciation, about 150,000 years ago), and an older glaciation of pre-Illinoian age (possibly as old as 2.5 million years ago). Hillslope and alluvial-fan and some of the stream sediments were laid down primarily under cold-climate conditions during the glacial periods. Wetland and flood-plain sediments are primarily of postglacial age. Weatheredbedrock materials are produced continuously under both cold and temperate climate by mechanical and chemical breakdown of bedrock.

The accompanying map and sections show the surface extent and subsurface relations of these deposits. Their composition and thickness, and the events they record, are described in the *Description of Map Units* (below). Well and boring data used to construct bedrock-surface elevation contours, and to infer the thickness and subsurface distribution of the deposits, are provided in tables 1 and 2 (in pamphlet). Figure 1 shows geomorphic and glacial features in the quadrangle, and figure 2 shows the thickness of surficial deposits and weathered bedrock. The chronologic relationships of the deposits and episodes of erosion are shown in the *Correlation of Map Units* (below). The hydrology of the surficial deposits and the history of geomorphic and glacial events in the quadrangle and adjacent areas are briefly described in the two following sections. Bedrock was mapped by Volkert (2008).

HYDROLOGY OF SURFICIAL DEPOSITS

Illinoian sand and gravel deposits (Qis) are productive glacial aquifers in valley fills in the Lamington River valley and the downstream reach of the Mill Brook valley. In these valleys, the sand and gravel is overlain by silt, clay, and fine sand of late Wisconsinan age (Qsl, Qdol; sections AA', CC'), which are confining or semiconfining beds. Sand and gravel beds in these valley fills are important regional aquifers and are tapped by public-supply wells to the north and west of the Mendham quadrangle (Canace and others, 1993; Schaefer and others, 1993; Nicholson and others, 1996). Only a few wells draw from the valley fills in the Mendham quadrangle (L1, L2, L3, L10, L12 in table 1). Elsewhere in the quadrangle, surficial deposits are generally too thin or insufficiently permeable to be aquifers, although several domestic wells draw from Illinoian sand and gravel beneath, or interbedded with, till in the moraine at the head of Shongum Lake (wells L18, L22), or from weathered bedrock beneath the moraine (well L20). A few other wells draw from granular weathered gneiss (Qwg, wells L26, 109) where it is sufficiently thick.

Hydraulic conductivities of surficial deposits may be estimated from aquifer-test and laboratory data on similar deposits in New Jersey (Stanford, 2000; Mennel and Canace, 2002). Sand and gravel deposits (units Qso, Qaf, Qis, Qst, and parts of Qal) are highly permeable, having estimated hydraulic conductivities that range from 10^1 to 10^3 feet per day (ft/d). Weathered rock, colluvium, and till with silty sand matrix (parts of Qwg, Qcg, and Qf) are also permeable, having estimated hydraulic conductivities from 10^{-1} to 10^2 ft/d. Glaciolacustrine silt and clay (Qsl, Qdol, Qpml, Qisl, and Qpsl) and weathered rock with clayey silt to silty clay matrix (parts of Qwcb, Qws, and Qwb) are of low permeability, having hydraulic conductivities of 10^{-5} to 10^{-3} ft/d. Fine sand and silt alluvial deposits (parts of units Qal, Qcal, and Qaf), silty to sandy silt till (parts of Qpt and Qf), and silty weathered rock and colluvium (parts of Qwcb, Qws, Qwb, Qwg, Qcg, and Qcs) are somewhat more permeable, having estimated hydraulic conductivities of 10^{-3} to 10^{-1} ft/d. Wetland deposits (Qs) and fill have variable hydraulic conductivities that depend on their clay and silt content. Peat with little mineral soil, and fill composed of sand, cinders, gravel, demolition debris, slag, and trash, may be highly permeable.

> **Gcal** unit Qcg, and alluvium consisting of dark brown to yellowish-brown or reddishyellow silty sand, sandy silt, to clayey silt, with some organic matter and beds and lag veneers of subangular to subrounded cobbles and boulders of gneiss, and rounded erratic cobbles and boulders of conglomerate and quartzite (adjacent to **figure 3 Photograph location** unit Qf). As much as 15 feet thick. Lag deposits are dominant in steeper reaches of valleys. Fine sediment, with organic matter, discontinuously overlies and infills lag deposits in gently sloping reaches. In some steep, narrow valleys, lags have moved downvalley to accumulate as bouldery lobes.

GEOMORPHIC AND GLACIAL HISTORY

Qsl GLACIAL LAKE SUCCASUNNA LAKE-BOTTOM DEPOSITS—Silt, clay, fine sand; light gray; laminated to thinly bedded. As much as 120 feet thick. In subsurface only (section AA'). Deposited in glacial Lake Succasunna. **Qr** ROCKAWAY RIVER OUTWASH—Sand and pebbly sand; light gray to very pale brown. As much as 20 feet thick. Deposited by meltwater flowing down the

GLACIAL LAKE PASSAIC, MOGGY HOLLOW STAGE, LAKE-BOTTOM **ODEROSITS—Silt, clay, very fine sand; light gray to very pale brown; laminated DEPOSITS—Silt, clay, very fine sand; light gray to very pale brown; laminated** to thinly bedded (fig. 3). As much as 70 feet thick. Deposited in Moggy Hollow stage of glacial Lake Passaic.

The highest elevations in the quadrangle are the plateau-like summit areas of uplands, including terrain above an elevation of about 850 feet in the northern half of the quadrangle and a smaller area between about 650 and 720 feet in elevation near the southern edge of the quadrangle (fig. 1). These uplands have about 150 feet of local relief and are as much as 700 feet above main valley floors. They are part of a regional low-relief upland erosion surface, termed the "Schooley peneplain" by Davis and Wood (1889), and "Kittatinny base level" by Salisbury (1898). They interpreted the feature as the product of fluvial erosion during a long period of stable base level. When base level lowered, the plain was incised by rivers, leaving upland remnants on resistant bedrock like the gneiss and granite in the Mendham quadrangle. Later studies in the Appalachian region proposed different origins for similar upland surfaces, and the peneplain hypothesis fell into disfavor. More recently, improved data on global sea level, and on the age of fluvial and marine sediments in New Jersey, indicate that a revised version of the peneplain hypothesis is plausible in this region. These data suggest that the upland erosion surface reached its final form in the middle to late Miocene (15-10 million years ago (Ma)) and was isolated on resistant-rock uplands by river incision in the late Miocene and early Pliocene (10-4 Ma) (Stanford and others, 2001). Erosion on moderate to gentle slopes has subsequently modified the surface, but it has done so at a rate much slower than that in the valleys, preserving its general form.

Qf FLANDERS TILL—Brown, brownish-yellow, and very pale brown silty sand $(5.10)(1.00)$ to sandy clayey silt with some $(5-10\%$ by volume) to many $(10-40\%)$ subrounded to subangular pebbles and cobbles and few (1-5%) to some subrounded boulders. Till matrix is generally compact, nonplastic, nonsticky, weakly jointed, and may have a weak to moderate subhorizontal fissility. Gravel consists chiefly of gray and white gneiss, some purple conglomerate and purple quartzite, and a little gray mudstone, sandstone, chert, and white-to-gray quartzite. Boulders are chiefly gneiss, purple conglomerate, and purple quartzite, with few to some white-to-gray quartzite. Gneiss clasts have weathered rinds generally more than 0.25-inch thick. Some mudstone and sandstone clasts are also weathered or fully decomposed. As much as 80 feet

Lowering sea level between 10 and 4 Ma due to growth of glaciers in Antarctica led to river incision along belts of erodible bedrock, including carbonate rock and micaceous, foliated gneiss. Carbonate rock includes Paleozoic dolomite in the Lamington River and North Branch of the Raritan River valleys, and marble in the Whippany River valley. Micaceous, foliated gneiss crops out in the Whippany River, Den Brook, and Mill Brook valleys. This incision lowered valley floors as much as 500 feet below the former base level on the Schooley surface.

silty sand, sandy clayey silt with some to many subangular gneiss pebbles and cobbles, in places underlain by, or interbedded with, thinly layered reddishyellow to pinkish-white clayey sand and sandy clay with few angular pebbles and cobbles. Long dimensions of clasts typically are aligned parallel to the hillslope. Upper blocky colluvium (fig. 4) is derived from downslope movement of fractured, weathered bedrock; lower, layered colluvium is derived from downslope movement of saprolite. Within the limit of pre-Wisconsinan glaciation (M1 on fig. 1), colluvium includes few to some erratic pebbles and cobbles of purple conglomerate, gray to white quartzite, and gray sandstone and mudstone from erosion of Flanders Till. Elsewhere, colluvium may include rare $(<0.1\%)$ conglomerate, quartzite and chert erratics from erosion of pre-Illinoian till. Colluvium on moderate-to-gentle slopes includes cobble-to-boulder lags formed by seepage erosion of weathered gneiss, alternating with, or overlain by, clayey-silty material. As much as 50 feet thick.

The pre-Illinoian glacier advanced into this landscape, perhaps as early as 2.5 Ma. It covered the entire quadrangle and advanced as far south as the Somerville area, 15 miles south of Mendham Borough. Magnetically reversed pre-Illinoian sediment (Ridge, 2004), weathering properties, erosional preservation, and pollen in basal lake sediment of probable pre-Illinoian age in Budd Lake (Harmon, 1968; Stanford and others, 2001), all indicate that the pre-Illinoian glaciation occurred in the early Pleistocene, between 2.5 and 0.8 Ma. Today, remnants of deeply weathered pre-Illinoian till (Qpt) crop out in only two small patches on flat uplands near Mendham Borough and Mount Freedom (Randolph Township). Additional deposits may be preserved beneath colluvium in similar topographic positions. Scattered erratic pebbles, cobbles, and boulders of purple conglomerate, gray to white quartzite, and gray chert (shown by symbols on the map) indicate that the pre-Illinoian till was formerly more widespread, but has been eroded from most of its original extent.

Qwb WEATHERED BASALT—Reddish-yellow to yellowish-brown clayey silt to $\frac{1}{2}$ silty clay with some to many angular to subangular pebbles and cobbles of basalt. Most clasts have reddish-yellow clayey-silty weathered rinds. As much as 15 feet thick.

Very fine sand, silt, and clay of possible pre-Illinoian glacial origin were observed under colluvium at two locations: 1) in a headwater valley in Mendham Borough and 2) along Primrose Brook in Jockey Hollow (both sites labeled "Qpsl" on map). A small deposit of deeply weathered sand and gravel (Qps) near Glen Alpin (Harding Township) may also be pre-Illinoian. These deposits may have been laid down in glacial lakes. The silt and fine sand at Mendham is in a small north-draining valley that would have been dammed by the glacier margin. The deposits at Glen Alpin and Primrose Brook may have been laid down in glacial Lake Watchung, a large ice-dammed lake that filled the upper Passaic River basin during the pre-Illinoian glaciation. More extensive deposits in this lake occur just south of the quadrangle, in the Bernardsville and Basking Ridge area (Stanford, 2008).

WEATHERED CARBONATE ROCK—Yellow, very pale brown, reddish-**Gwcb** vertically clayey silt to silty clay, minor sandy silt, with some to many yellow, light gray clayey silt to silty clay, minor sandy silt, with some to many light gray, yellow, and reddish-yellow angular chips and pebbles of carbonate rock. As much as 100 feet thick but thickness is highly variable. Crops out in one bank along India Brook; present beneath colluvium in the North Branch Raritan River valley and beneath glacial deposits in the Lamington valley

- **Quarry or mine—**Inactive in 2009.
- **L15** Well with log in table 1—Location accurate within 100 feet. Wells located from tax-parcel maps, street addresses, or distances from landmarks as reported by drillers.
- **L22 Well with log in table 1—Location accurate within 500 feet.**
- **Well reporting thickness of surficial material—**Data in table 2**.** Location **433** ! accurate within 100 feet. **Well reporting thickness of surficial material—**Data in table 2**.** Location **427** .
- accurate within 500 feet. Qis • Subsurface unit exposed—Observed in excavation or streambank, beneath mapped surface unit. "Qcgo" indicates weathered colluvium observed beneath fresh colluvium. "Qpmb" indicates beach sand and gravel of the Moggy Hollow stage of Lake Passaic. "Qpsl" and "Qisl" indicate glaciolacustrine silt, clay, and very fine sand of pre-Illinoian and Illinoian
- age, respectively. "Qis/Qisl" indicates Illinoian silt and clay under Illinoian sand and gravel. \Diamond **Pre-Illinoian erratic**—One or more cobbles and boulders of purple conglomerate, white to gray quartzite, or gray chert, from pre-Illinoian glaciation.
- Glacial ridge—Line on crest, barbs point toward glacier. Emplaced along Illinoian glacial margins.
- **Seepage scarp**—Line at position of groundwater emergence. Water drains downslope from this position. Seepage is also common along upland margins of units Qal and Qcal.
- **Elevation of bedrock surface—Contour interval 50 feet. Includes top** surface of weathered bedrock. Shown only where glacial deposits are more than 25 feet thick.
- **Bedrock outcrop—**Outcrops within unit Qwgt are not shown, and some small outcrops along streams or in artificial cuts are also not shown. Refer to Volkert (2008) for these outcrop locations.
- **Lake or pond—Obscures underlying deposits.**

A long period of weathering and erosion followed the pre-Illinoian glaciation. Almost all of the pre-Illinoian deposits were eroded away, and rivers deepened their valleys by several tens of feet below the pre-Illinoian land surface.

A second glacier entered this landscape during the middle Pleistocene, probably during the late Illinoian glaciation about 150,000 years ago. It advanced to position M1 (fig. 1). This glaciation cannot be directly dated, but its deposits are much less weathered and eroded than those of the pre-Illinoian, indicating a significantly more recent age. It is older than late Wisconsinan, as indicated by a radiocarbon date of 36,070+/-280 yrs. BP (41,650-40,840 calibrated yrs. BP) (Beta 309763) on peat overlying Illinoian sand and gravel, and overlain by six feet of colluvium, at Dalrymple Pond (latitude 40°51'09", longitude 74°34'50"). Deposits of this glacier include till (Flanders Till of Stone and others, 2002, unit Qf), sand and gravel (Qis), and silt and clay (Qisl). The sand, gravel, silt, and clay were laid down by meltwater in glacial lakes. The distribution of these sediments shows that lakes formed in the Lamington River, Jackson Brook, Wallace Brook, Mill Brook, and Den Brook valleys. These valleys all drained northward and so were dammed by the glacier to create lakes.

Before it was filled with glacial sediment, the upper Lamington River valley drained to the north as a tributary to the Rockaway River (Stanford, 1989). During the Illinoian glaciation, this valley was dammed by the ice front to form glacial Lake Ironia (Stone and others, 2002). Lake Ironia was controlled by a spillway draining southward across the former drainage divide into the Raritan basin at an elevation of about 680 feet near Chester Borough (5 miles west of Mendham Borough). The northwest corner of the quadrangle includes a small part of this lake. The lake drained when the retreating ice front uncovered the Rockaway River valley at Dover, permitting eastward flow down the Rockaway.

Smaller lakes occupied the Jackson, Wallace, Mill, and Den Brook valleys, all of which are also north-draining tributaries to the Rockaway River (fig. 1). In the Jackson and Wallace Brook valley, high and intermediate lake stages (JB1 and JB2, respectively, on fig. 1) were controlled by spillways at 950 and 870 feet in elevation that drained west into the Lamington River valley. A series of temporary, lowering spillways operated during the drop from the high to intermediate stage. The intermediate stage was succeeded by a low stage (JB3 on fig. 1) controlled by a spillway at 840 feet that drained southward into the Mill Brook valley. The low stage drained as ice uncovered the Rockaway valley at Dover, about 0.5 mile north of the quadrangle boundary.

The Den and Mill Brook valleys were occupied by glacial Lake Shongum. The high stage of Lake Shongum (SH1 on fig. 1) was controlled by a spillway on the Den Brook drainage divide at an elevation of 850 feet that drained southward into the Whippany River valley. The high stage extended westward into the Mill Brook valley when the ice front retreated from a gap west of Shongum Lake. Before this retreat, a small lake was ponded in the headwaters of Mill Brook south of the present site of Dalrymple Pond (DP on fig. 1). This lake drained along the ice front banked against the hillslope to the east. The high stage of glacial Lake Shongum lowered to the intermediate stage (SH2 on fig.1) when the rereating ice front uncovered a lower gap at an elevation of 800 feet just south of the present alignment of Route 10, which allowed eastward drainage into the Whippany valley. Shortly thereafter, continued retreat exposed a gap at an elevation of 700 feet just to the north, along the present Route 10, and the lake lowered to its low stage (SH3 on fig. 1). Both the intermediate and low stages extended into the Mill Brook valley. The low stage lowered again and then drained as the retreating ice front uncovered the Rockaway valley north of Mount Tabor, about 1.5 miles northeast of the northeast corner of the quadrangle. Illinoian glacial deposition in the Mendham quadrangle ended at this time.

In glacial lakes Ironia and Shongum, sand and gravel were deposited in lacustrine fans and deltas, and small volumes of silt and clay (symboled "Qisl" where observed) were deposited on lake bottoms, during both advance and retreat of the glacier. At and north from the terminal position (M1 on fig. 1), advancing ice overrode the advance-phase lake sediments and emplaced till atop them. This overriding action built a moraine in the Den and Mill Brook valleys (between margins M1 and M2 on fig. 1). This moraine is especially prominent at the north end of Shongum Lake, where it is more than 150 feet thick. Till overlying lacustrine deposits was observed in a pipeline trench here and in the Lamington valley, and in a former gravel pit west of Openaka Lake (symboled on map).

A period of weathering, and stream and hillslope erosion, followed the Illinoian glaciation. During periods of cold climate, primarily during the early and late Wisconsinan, forest cover was reduced and permafrost developed, impeding soil drainage and thereby waterlogging the surficial material during thaws. Weathered-rock material and glacial sediment on hillslopes became unstable and moved downslope to form aprons of colluvium (Qcg, Qcs). Where the material was transported downslope into steep tributary channels, streams flushed it into main valleys to form alluvial fans (Qaf). North of the Illinoian limit, till on moderate-to-steep hillsides was eroded away, exposing the underlying weathered rock. Along footslopes, particularly in the Mill Brook valley, colluvium accumulated on top of lacustrine deposits and till (sections AA', CC'). Exposures where lacustrine sediment was observed beneath colluvium are indicated by symbols on the map.

Most colluvium at the surface is lightly weathered and probably of Wisconsinan age. The 36,070+/-280 yrs. BP date on peat beneath six feet of colluvium at Dalrymple Pond (cited above) indicates that colluvium here was deposited during the late Wisconsinan. South of the Illinoian limit, older, weathered colluvium, deposited during earlier periods of cold climate, likely occurs in the subsurface of thick aprons. Older colluvium containing weathered and decomposed gneiss clasts, and showing soil development, was observed in two exposures (symboled with "Qcgo' on map) beneath fresh-clast colluvium.

The most-recent glacier, known as the late Wisconsinan, reached its southernmost position by about 25,000 years ago (21,000 radiocarbon years ago), based on radiocarbon dates of pre-advance organic material in a sediment core from Budd Lake (Harmon, 1968; Stanford and Witte, 2002) and on Long Island, New York (Sirkin and Stuckenrath, 1980). The glacier advanced to within two miles of the northern and eastern edge of the Mendham quadrangle, but did not enter the map area. At its maximum position, and during retreat, this glacier impounded lakes in the Lamington valley (glacial Lake Succasunna, SC on fig. 1), in the Rockaway valley (glacial Lake Dover, DO on fig. 1), and in the Passaic valley (Moggy Hollow stage of glacial Lake Passaic, MH on fig. 1).

In Lake Succasunna, which was similar to Illinoian glacial Lake Ironia, a large delta filled the northern reach of the lake. This delta was built outward from the terminal position (about one mile north of the northwest corner of the quadrangle) and consists of sand and gravel (Qso) topset and foreset beds overlying silt, fine sand, and clay bottomset and lake-bottom beds (Qsl on section AA').

In the Mendham quadrangle, deposition in Lake Dover was exclusively lakebottom silt, fine sand, and clay (Qdol on section CC'). Glaciofluvial sand and gravel (Qr) was laid down atop the Lake Dover sediments by meltwater draining down the Rockaway valley after Lake Dover drained (Stanford, 1989).

Lake Passaic was a large lake that filled the upper Passaic basin and had three glacial-phase stages (Stone and others, 2002). Only the highest, known as the Moggy Hollow stage, extended into the Mendham quadrangle. It was controlled by a spillway on the Passaic-Raritan drainage divide near Far Hills, 7 miles south of Mendham Borough. In the quadrangle, it flooded the Whippany and Passaic valleys to an elevation of 350 to 360 feet. Silt, clay, and fine sand (Qpml) accumulated on the lake bottom (fig. 3). In the Whippany valley, colluvium accumulating on footslopes along the lakeshore became interbedded with lake clay. Places where lake clay was observed within colluvium are symboled on the map.

A cut along Interstate Route 287 near Glen Alpin (Harding Township) exposed a thin deposit of rounded gneiss gravel with a few quartzite, chert, and siltstone erratics, and laminated to cross-bedded coarse sand, within colluvium at an elevation of about 350 feet (symboled "Qpmb" on map). Because there is no fluvial source for this material, it is most likely a beach deposit, with ice-rafted erratics, deposited by wave action on the shoreline of the Moggy Hollow stage.

After the Moggy Hollow stage lowered, sand and gravel (Qst) was laid down atop the lake clays in places by the Whippany River and its tributaries. Sand and gravel under colluvium along Primrose Brook near Glen Alpin (symboled on map) are also post-Moggy Hollow stream deposits. These deposits are between 300 and 340 feet in elevation, and may be graded to the Great Notch stage of Lake Passaic, which was at 290 to 300 feet in elevation in this area (Stanford, 2006). Today, these deposits form terraces several feet above the modern flood plain.

Radiocarbon dates on postglacial organic material in New Jersey and adjacent areas indicate that glacial lakes had drained, and late Wisconsinan glacial deposition in the Mendham quadrangle had ended, by 19,000 radiocarbon years ago (Stone and others, 2002). As climate warmed, forest cover returned and hillslopes stabilized. Deposition of colluvium and alluvial fans ceased and streams and groundwater seepage eroded and incised the colluvial aprons and alluvial fans in places. The seepage and stream erosion formed cobble-toboulder lags and flood plains (Qcal, Qal). Seepage and slopewash also laid down clayey silt at the foot of some gently sloping colluvial aprons. A radiocarbon date of 1290+/-30 yrs. BP (1290-1170 calibrated yrs. BP) (Beta 309792) on plant material under four feet of clayey silt at one such site near Mendham (latitude $40^{\circ}46'16''$, longitude $74^{\circ}34'47''$) indicates that the silt here accumulated within the past 1300 years, perhaps chiefly from agricultural clearing during European colonization. Wetland sediment (Qs) accumulated in poorly drained areas. Flood-plain and wetland deposition continues today.

DESCRIPTION OF MAP UNITS

Postglacial Deposits—These include man-made fill, stream deposits in fans (Qaf), modern channels and flood plains (Qal, Qcal), and low terraces (Qst), and wetland deposits in swamps and marshes (Qs).

ALLUVIAL FAN DEPOSITS—Pebble-to-cobble gravel, cobble-to-boulder **Qaf** gravel, sand, silt; brown, yellowish-brown, gray; moderately sorted, stratified. As much as 25 feet thick (estimated). Some small alluvial fans are included with units Qal and Qcal.

Glacial Deposits—These include till and meltwater sediments laid down during the late Wisconsinan, Illinoian, and pre-Illinoian glaciations. Till is a poorly sorted, nonstratified sediment containing gravel and boulders, deposited directly from glacial ice (units Qf and Qpt). The meltwater sediments are moderately to well-sorted and stratified and include sand and gravel laid down in glacial lakes (Qso, Qis, Qps) and glaciofluvial plains (Qr), and silt, clay, and fine sand laid down in glacial lakes (Qsl, Qdol, Qpml).

Deposits of late Wisconsinan age GLACIAL LAKE SUCCASUNNA DELTAIC DEPOSITS—Fine-to-medium **Qso** sand, minor coarse sand and pebbly sand; very pale brown to light gray. As much as 40 feet thick. Deposited in glacial Lake Succasunna. Extensively mined in quadrangle, may have been mostly removed from mined areas.

Deposits of Illinoian age

thick.

ILLINOIAN STRATIFIED DEPOSITS—Pebble-to-cobble gravel, fine-tocoarse sand, silty sand, minor cobble-to-boulder gravel. Sand and silty sand are yellowish-brown to brown. Gravel consists chiefly of gray and white gneiss, some purple conglomerate and purple quartzite, and minor gray sandstone and mudstone, quartzite, and chert. Gneiss clasts, and some sandstone and mudstone clasts, have weathered rinds or are partially decomposed. Sand is unstratified to plane- and cross-bedded. As much as 130 feet thick. Deposited in glacial lakes in the Lamington, Jackson, Mill, and Den Brook valleys. Commonly interbedded or overlain by Flanders Till (Qf).

Deposits of pre-Illinoian age

PRE-ILLINOIAN TILL—Yellowish-brown to reddish-yellow sandy silt to sandy clayey silt, with some subrounded to subangular pebbles and cobbles, and few subrounded boulders. Gravel consists chiefly of gray and white gneiss, some purple conglomerate and quartzite, and a little gray to white quartzite, gray mudstone and sandstone, and dark-gray chert. Boulders are chiefly gray gneiss, purple conglomerate, and gray to white quartzite. The mudstone, sandstone, and gneiss gravel clasts have weathered rinds or are completely decomposed. As much as 10 feet thick. Equivalent to the Port Murray Formation, till facies, of Stone and others (2002).

PRE-ILLINOIAN STRATIFIED DEPOSIT—Reddish-yellow to yellowishbrown clayey silty sand to clayey silt with some to many pebbles and fine cobbles. Gravel consists chiefly of gray and white gneiss, some red shale and siltstone, and a little gray, white, and red quartzite. Gneiss, shale, and siltstone clasts are deeply weathered to fully decomposed. As much as 20 feet thick (estimated). In erosional remnant of a lacustrine deposit near Glen Alpin (Harding Township).

Hillslope Deposits—Nonstratified, poorly sorted sediment deposited at the foot of hillslopes by downslope movement of material. GNEISS COLLUVIUM—Yellowish-brown, reddish-yellow, brown sandy silt,

SILTSTONE AND CONGLOMERATE COLLUVIUM—Brown, yellowish**brown, light gray sandy clayey silt with some subangular pebbles and cobbles.** Gravel is weathered from conglomerate bedrock and consists chiefly of gray quartzite and sandstone. As much as 10 feet thick.

Weathered Bedrock Material—Nonstratified, poorly sorted sediment and fractured rock formed by mechanical and chemical weathering of bedrock.

WEATHERED SILTSTONE, SHALE, AND CONGLOMERATE—Reddish-**EXAMPLE SHATHER STATE, AND CONCLOMENTE REGISTER**
brown, yellowish-brown, brown clayey silt to silty clay with some angular to subangular gray to red siltstone and shale chips and some gray and purple quartzite pebbles and cobbles. As much as 30 feet thick.

ARTIFICIAL FILL—Artificially emplaced sand, gravel, silt, clay, and rock fragments, and man-made materials including cinders, ash, brick, concrete, wood, slag, asphalt, metal, glass, and trash. Color varied but generally dark \mathbb{Z} brown, gray, or black. In highway and railroad fills, dams, dikes, made land, waste-rock disposal piles, and trash fills (aft). As much as 50 feet thick, generally less than 20 feet thick. Many small areas of fill are not mapped. aft

ALLUVIUM—Sand, silt, pebble-to-cobble gravel, cobble-to-boulder gravel, minor clay and peat; fine sediment is dark brown, brown, yellowish-brown, gray; moderately to well-sorted, stratified. Contains varied amounts of organic matter. Sand and gravel are deposited in active channels. Sand, silt, clay and peat are deposited in back channels, overbank areas, and groundwater seepage areas, chiefly on broad flood plains. In places, alluvium includes cobble-to-boulder lags from erosion of weathered gneiss and colluvium. As much as 15 feet thick. **Qal**

WEATHERED GNEISS—Yellowish-brown, yellow, very pale brown, reddishyellow, silty sand, silty clayey sand to sandy clayey silt, locally micaceous, with **Qwg** few to many subangular pebbles and cobbles of gneiss. Includes mixed clast**a** and-matrix sediment, granular decomposed rock, fractured rock rubble, and and-matrix sediment, granular decomposed rock, fractured rock rubble, and saprolite that preserves original rock structure (fig. 5). Clasts vary from unweathered to fully decomposed. On gentle to moderate slopes, well records indicate that clast-and-matrix sediment (described by drillers as "overburden", "hardpan", "sandy hardpan", and "clay hardpan"), which is fractured rock mixed with sandy-clayey saprolitic material by colluviation, cryoturbation, and bioturbation, is generally between 5 and 30 feet thick and commonly overlies or grades downward to saprolite (described by drillers as "rotten rock", "sandstone", "rotten granite", and "soft granite") that may be as much as 100 feet thick over unweathered rock. On steep slopes, fractured-rock rubble, generally less than 20 feet thick, overlies unweathered bedrock. Total thickness of weathered material is as much as 150 feet but is generally less than 50 feet (fig. 2). The uppermost, clast-and-matrix material may contain traces of quartzite, chert, and gray sandstone and mudstone erratic pebbles and cobbles, especially within the pre-Wisconsinan glacial limit. "Qwgt" indicates areas, typically on the steepest slopes and narrow ridgetops, where weathered material is thin or absent and fractured outcrop abundant.

STREAM TERRACE DEPOSITS—Fine-to-medium sand, silt, pebble-to-cobble gravel; brown, very pale brown, yellowish-brown, light gray; moderately to well sorted, stratified. As much as 10 feet thick. Form terraces with surfaces 3-10 feet above the modern flood plain. **Qst**

Qs SWAMP AND MARSH DEPOSITS—Peat and organic silt, clay, and minor fine \Box sand; black, dark brown, and gray. As much as 10 feet thick (estimated). ALLUVIUM AND COLLUVIUM, UNDIVIDED—Interbedded colluvium as in (section AA').

MAP SYMBOLS

Radiocarbon date—In radiocarbon years before present, with error and laboratory number.

Excavation perimeter—Line encloses excavated area. Outlines pits and

quarries.

1290+/-30 ●
(Beta 309762)

Sand and gravel pit—Inactive in 2009.

RRELATION OF MAP UNITS

REFERENCES

Contact—Long-dashed where approximately located, short-dashed where gradational or feather-edged, dotted where removed by excavation. **Material observed in hand-auger hole, exposure, or excavation** !

- Canace, Robert, Stanford, S. D., Hall, D. W., 1993, Hydrogeologic framework of the middle and lower Rockaway River basin, Morris County, New Jersey: N. J. Geological Survey Report GSR 33, 68 p. Davis, W. M., and Wood, J. W., 1889, The geographic development of northern New Jersey: Proceedings of the Boston Society of Natural History, v. 24, p.
- 365-423. Harmon, K. P., 1968, Late Pleistocene forest succession in northern New Jersey: New Brunswick, N. J., Rutgers University, Ph. D. dissertation, 203 p. Mennel, W. J., and Canace, Robert, 2002, New Jersey Geological Survey hydro database: N. J. Geological Survey Digital Geodata Series DGS 02-1,
- www.state.nj.us/dep/njgs/geodata/dgs02-1.zip Nicholson, R. S., McAuley, S. D., Barringer, J. A., Gordon, A. D., 1996, Hydrogeology of, and ground-water flow in, a valley-fill and carbonate-
- rock aquifer system near Long Valley in the New Jersey Highlands: U. S. Geological Survey Water Resources Investigations 97-4157, 159 p. Ridge, J. C., 2004, The Quaternary glaciation of western New England with correlations to surrounding areas, *in* Ehlers, J., and Gibbard, P. L., eds., Quaternary glaciations—extent and chronology, part II: Elsevier, p. 169-
- 199. Salisbury, R. D., 1898, The physical geography of New Jersey: N. J. Geological Survey Final Report of the State Geologist, v. 4, 200 p. Schaefer, F. L., Harte, P. T., Smith, J. A., and Kurtz, B. A., 1993, Hydrologic conditions in the upper Rockaway River basin, New Jersey, 1984-86: U. S.
- Geological Survey Water Resources Investigations Report 91-4196, 103 p. Sirkin, L. A., and Stuckenrath, R., 1980, The Portwashingtonian warm interval in the northern Atlantic Coastal Plain: Geological Society of America
- Bulletin, v. 91, p. 332-336. Stanford, S. D., 1989, Surficial geology of the Dover quadrangle, Morris and Sussex counties, New Jersey: N. J. Geological Survey Geologic Map Series GMS 89-2, scale 1:24,000.
- Stanford, S. D., 2000, Glacial aquifers of New Jersey, *in* Harper, D. P., and Goldstein, F. R., eds., Glacial geology of New Jersey: field guide and proceedings for the seventeenth annual meeting of the Geological Association of New Jersey: Trenton, N. J., Geological Association of New
- Jersey, p. IV.1-IV.21. Stanford, S. D., 2006, Surficial geology of the Morristown quadrangle, Essex and Morris counties, New Jersey: N. J. Geological Survey Open-File Map OFM 67, scale 1:24,000. Stanford, S. D., 2008, Surficial geology of the Bernardsville quadrangle, Morris
- and Somerset counties, New Jersey: N. J. Geological Survey Open-File Map OFM 74, scale 1:24,000. Stanford, S. D., Ashley, G. M., and Brenner, G. J., 2001, Late Cenozoic fluvial stratigraphy of the New Jersey Piedmont: a record of glacioeustasy,
- planation, and incision on a low-relief passive margin: Journal of Geology, v. 109, p. 265-276. Stanford, S. D., and Witte, R. W., 2002, Surficial geology of the Tranquility quadrangle, Warren, Sussex, and Morris counties, New Jersey: N. J. Geological Survey Open-File Map OFM 51, scale 1:24,000. Stone, B. D., Stanford, S. D., and Witte, R. W., 2002, Surficial geologic map of northern New Jersey: U. S. Geological Survey Miscellaneous
- Investigations Map I-2540-C, scale 1:100,000. Volkert, R. A., 2008, Bedrock geologic map of the Mendham quadrangle, Morris and Somerset counties, New Jersey: unpublished map, N. J. Geological Survey, scale 1:24,000.

Well on sections—Projected to line of section. Owing to projection, depths of contacts on section may not be identical to those in well. **127**

Geology mapped 1988, 2010-2011 Cartography by S. Stanford

74o30' 40o'45'

40o52'30"

50'

MORRISTOWN

 \mathcal{G} δ

47'30"

SURFICIAL GEOLOGY OF THE MENDHAM QUADRANGLE MORRIS AND SOMERSET COUNTIES, NEW JERSEY OPEN-FILE MAP OFM 94

Pamphlet containing tables 1 and 2 accompanies map

DEPARTMENT OF ENVIRONMENTAL PROTECTION WATER RESOURCE MANAGEMENT NEW JERSEY GEOLOGICAL AND WATER SURVEY

Qis

Qpt

Qps

Qcg

Qcs

r

summer. Each pair of clay and silt beds represents a year of deposition and is known as a varve. Location shown on inset and map.

> Figure 4. Blocky facies of gneiss colluvium. Note that long dimensions of clasts are

Figure 5. Weathered gneiss, showing angular fractured-rock rubble and coarse sandy saprolite. Location shown on inset and map.

Surficial Geology of the Mendham Quadrangle Morris and Somerset Counties, New Jersey

New Jersey Geological Survey Open-File Map OFM 94 2012

pamphlet with tables 1 and 2 to accompany map

Table 1.--Selected well and boring records.

¹ Identifiers consisting of 4 or 5 digits are well permit numbers issued by the N. J. Department of Environmental Protection, Bureau of Water Allocation. All are prefixed with "25-". Identifiers of the form "Nxx-xx-xxx" are N. J. Atlas Sheet coordinates of well logs in the N. J. Geological Survey permanent note collection.

 2 Depth (in feet below land surface) and driller's or logger's description is provided. Notation "NR" indicates depth of contact is not reported. Inferred map units and author's comments are indicated in parentheses. All descriptions are reproduced as they appear in the original source, except for minor format, punctuation, and spelling changes. Logs identified as "abbreviated" have been condensed for brevity, or have minor details omitted. Many bedrock descriptions have been condensed; these are not identified as abbreviated. Map units are inferred from the known extent of materials at the surface and from known depositional settings, in addition to the drillers' descriptions. For wells completed in surficial material, the yield (in gallons per minute, gpm) and depth to which the well is cased, if reported, are included after the description.

Table 2.—Selected wells and borings reporting thickness of surficial material. All depths are in feet below land surface.

Well Number	Identifier ¹	Depth (ft.) to	Depth (ft.) to	Total depth
		base of	base of	(f _t)
		surficial	saprolite ³	
		material ²		
$\mathbf{1}$	32251	18		150
$\sqrt{2}$	21636	$\overline{4}$	16	200
$\overline{\mathbf{3}}$	Healy 1961, B9	11		12
$\overline{\mathbf{4}}$				12
	Healy 1961, B3	10		
$\overline{5}$	8857	$\overline{33}$		82
$\overline{6}$	27061	50		100
$\overline{7}$	30173	40		$70\,$
$\bf 8$	26711	30		230
9	34110	8		200
10	40231	18	43	130
$\overline{11}$	20219	21		173
12	23103	13	45	95
13	34534	85		275
14	20220	8	75	150
$\overline{15}$	27586	15		225
16	22360	40	$70\,$	150
17	17853	10		400
18	26332	$10\,$	35	500
19	21567	12	$22\,$	123
$20\,$	36114	6		400
$\overline{21}$	23819	80	89	198
22	23820	111		223
23	15125	35		97
$24\,$	13360	6		216
$\overline{25}$	17878	15		123
$26\,$	26343	40	60	100
27	33719	$11\,$	31	122
28	3823	18		$\overline{55}$
29	4044	5	15	48
30	8408	$\overline{26}$		80
$\overline{31}$	3824	$\overline{50}$	>65	
32	19520	60		122
$\overline{33}$	6712	26		78
34	8884	35	52	62
$\overline{35}$	12663	$\overline{4}$		122
36	8546	13		60
37	21154	56		98
38	22601	40		147
39	22834	61		147
40	21754	$\bf 8$		225
41		$\overline{4}$		
	42003			250
42	27819	18		180 214
43	14610	36		
$44\,$	41290	>70		
45	41188	$>\!\!40$		
$46\,$	10362	10		178
47	3926	6		300
48	5596	14		377
49	19470	$25\,$	60	140
50	34083	$\overline{2}$		250
$51\,$	24249	$\overline{2}$	30	225
52	30290	100		330
53	12387	11	75	121
54	12481	18		124

¹Identifiers consisting of 3, 4, or 5 digits are well permit numbers issued by the N. J. Department of Environmental Protection. All are preceded by "25-". Identifiers of the form 25-x-xxx are N. J. Atlas Sheet coordinates of well logs in the N. J. Geological Survey permanent note collection. Identifiers of the form "Bx Healy xxxx" and "Bx Wash. Valley" are test borings on file at the N. J. Geological Survey.

 2 Described by drillers as "overburden", "hardpan", "clay and gravel", "clay", "sand and clay", "sandy hardpan", and "stony hardpan".

³Described by drillers as "sandstone", "rotten rock", "rotten granite", "soft granite", "broken rock", "fractured rock", "weathered rock", and "soft sandstone".